

What controls the explosivity of subglacial eruptions?

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Torfajökull









Subglacial Torfajökull eruptions



- \diamond Mixed rhyolite-tholeiite
- Alkali basalt

 $\langle \rangle$







Bláhnúkur





Jakobsson and Guðmundsson (2008), Jökull















































Angel Peak





Explosivity vs ice thickness

Edifice	Part of ring fracture unit?	Volume (km³)	Eruptive environment	Inferred ice thickness (m)*	Inferred eruptive style
Angel Peak	Yes	<0.1	Subglacial	120ª	Effusive
Bláhnúkur	No	<0.1 ^b	Subglacial⁵	400 ^c	Effusive ^b
Dalakvísl	Yes	<0.2 ^g	Subglacial ^{†g}	330ª	Mixed: effusive-explosive ^g
SE Rauðfossafjöll	Yes	~1 ^h	Emergent ^h	290 ⁱ	Explosive ^h
NW Rauðfossafjöll	Yes	~1	Emergent	290 ⁱ	Explosive

Owen et al., (2013), Geology

No relationship

Dalakvísl





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Dalakvísl obsidian sheets



Owen et al., (in prep.), Jökull



Dalakvísl obsidian sheets



Owen et al., (2013), JVGR

 $1000\ \mu m$

Dalakvísl







Dalakvísl Pre-eruptive volatile content



Effusive lava lobes: 2.4 wt.% H₂O





More explosive obsidian sheets: $4.8 \text{ wt.\% H}_2\text{O}$



Dalakvísl degassing path





Dalakvísl





Torfajökull pre-eruptive volatile content





Torfajökull degassing path





Torfajökull degassing path



Owen et al., (2013), Geology



Torfajökull conclusion

The behaviour of subglacial rhyolite

- Is little affected by the quantity of ice loading
- However rapid decompression may trigger a transition to more explosive activity
- There is also a correlation between explosivity and
 - Pre-eruptive volatile content
 - Degassing path

Basalt in Iceland





Eyjafjallajökull 2010





Grímsvötn 2011



Holuhraun 2014



Katla




Katla



Katla



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Eyjafjallajökull 2010

Grímsvötn 2011



Katla1918

Katla vs Eyjafjallajökull

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	Katla 1918	Eyjafjallajökull 2010
Date of commencement	12 th Oct 1918 ^A	14 th Apr 2010 ^G
Duration of eruption	24 days ^A	39 days ^G
Composition	Basalt (47% SiO ₂) ^B	Benmoreite and trachyte ^H
VEI	4 (at least) ^c	4 (upgraded from 3) ^c
Total erupted volume (DRE)	1 km ^{3 D}	0.2 km ^{3 H}
Max plume height	14 km ^A	10 km ^H
Volume of airborne tephra	0.7 km ^{3 D}	<0.3 km ^{3 H}
Area of tephra fall on land	50,000 km ^{2 A}	12,000 km ²¹
Thickness of ice over eruption site	400 m ^D	200 m ^H
Volume of subglacial lavas	0.2 km ^{3 E}	0.02 km ^{3 H}
Time taken to melt overlying ice	2 hours ^E	3-4 hours ^H
Jökulhlaup volume	>8 km ^{3 E}	<0.06 km ³¹
Flooded area	600-800 km ^{2 F}	57.5 km ²¹
Max discharge rate of jökulhlaup	>300,000 m ³ s ^{-1 E}	2,600 m ³ s ⁻¹¹
Volume of flood transported tephra	0.7-1.6 km ^{3 F}	0.03 km ^{3 H}

A: Larsen (2010); B: Óladóttir et al., (2008); GVP (2013); D: Sturkell et al., (2010); E: Tómasson, (1996); F: Larsen (2000); H: Guðmundsson et al., (2012); I: Gylfason et al., (2012)



























Katla 1918 air fall tephra



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Katla1918 air fall tephra





Katla1918 air fall tephra





Katla1918 air fall tephra



Katla 1918 Grain –size distributions



Guðmundsson et al., (2012), Sci Reports

Lancaster 558 University

Environment

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Katla 1918 Grain –size distributions





Katla 1918 FTIR: air fall deposits



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0.07 wt.%









- 130 m of meltwater
- 120 m of ice (full thickness ~400 m)
- 40 m of rock (e.g. if fragmentation occurred within the conduit)
- hydration



Katla 1918 TGA





Katla 1918 textures – bubble size





Katla 1918 textures - shearing







Katla 1918 Textures – coalescence





Katla 1918 Textures – connectivity



100 µm



Katla 1918 Hostage – coalescence





Katla 1918 Textures – coalescence





Katla 1918 Textures - microlites





Katla 1918 Textures – bubble collapse



Katla 1918 Textures - welding

1 mm



Katla 1918 textures - mingling









1 mm



- Vesiculation
- Fragmentation
- Bubble collapse
- Welding
- Vesiculation
- Fragmenation



1 mm

Katla 1918 hotstage

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- Torfajökull the volatiles did it!
- Katla the volatiles might've done it.... Lots of further work required
- Lots of evidence of repeated fragmentation (based on welded clasts) possibly with additional vesiculation and/or re-melting between the fragmentation events... so did fragmentation occur in the conduit????
- Some evidence that the jökulhlaup samples quenched rapidly in a water-rich environment under elevated pressure whereas the air-fall tephra cooled more slowly (high microlite content and larger bubbles in clast center) under atmospheric conditions (FTIR)
Further work



- Detailed investigation of the different layers in the air-fall tephra collected this summer – (imaging, FTIR, TGA)
- Characterisation of external grain morphology
- SEM vesicle size distribution and bubble number densities
- Chemical analysis: LA-ICP-MS for trace elements and EPMA for major elements and some volatiles
- Do the different grain sizes in the air-fall tephra layers represent different extents of fragmentation?
- And if so, what was causing the different eruptive behaviour???

Any questions?



