

1 **Ground-based observations of Saturn’s auroral**
2 **ionosphere over three days: trends in H_3^+ temperature,**
3 **density and emission with Saturn local time and**
4 **planetary period oscillation**

5 James O’Donoghue^{d,a}, Henrik Melin^a, Tom S. Stallard^a, G. Provan^a,
6 Luke Moore^d, Sarah V. Badman^e, Stan W.H. Cowley^a, Kevin H. Baines^c,
7 Steve Miller^b, and James S.D. Blake^a

8 ^a*Department of Physics and Astronomy, University of Leicester, Leicester, LE1 7RH, UK*

9 ^b*Atmospheric Physics Laboratory, Department of Physics and Astronomy, University College*
10 *London, London, WC1E 6BT, UK*

11 ^c*NASA Jet Propulsion Laboratory, M/S 183-601, 4800 Oak Grove Drive, Pasadena, CA 91109,*
12 *USA*

13 ^d*Center for Space Physics, Boston University, Boston, MA 02215, USA*

14 ^e*Department of Physics, Lancaster University, UK*

15 *Key words:* Saturn, aurora, magnetosphere, ionosphere, atmosphere

Abstract

On 19 to 21 April 2013, the ground-based 10-metre W.M. Keck II telescope was used to simultaneously measure H_3^+ emissions from four regions of Saturn's auroral ionosphere: 1) the northern noon region of the main auroral oval; 2) the northern midnight main oval; 3) the northern polar cap and 4) the southern noon main oval. The H_3^+ emission from these regions was captured in the form of high resolution spectral images as the planet rotated. The results herein contain twenty-three H_3^+ temperatures, column densities and total emissions located in the aforementioned regions - ninety-two data points in total, spread over timescales of both hours and days. Thermospheric temperatures in the spring-time northern main oval are found to be cooler than their autumn-time southern counterparts by tens of K, consistent with the hypothesis that the total thermospheric heating rate is inversely proportional to magnetic field strength. The main oval H_3^+ density and emission is lower at northern midnight than it is at noon, in agreement with a nearby peak in the electron influx in the post-dawn sector and a minimum flux at midnight. Finally, when arranging the northern main oval H_3^+ parameters as a function of the oscillation period seen in Saturn's magnetic field - the planetary period oscillation (PPO) phase - we see a large peak in H_3^+ density and emission at $\sim 115^\circ$ northern phase, with a full-width at half-maximum (FWHM) of $\sim 44^\circ$. This seems to indicate that the influx of electrons associated with the PPO phase at 90° is responsible at least in part for the behavior of all H_3^+ parameters. A combination of the H_3^+ production and loss timescales and the $\pm 10^\circ$ uncertainty in the location of a given PPO phase are likely, at least in part, to be responsible for the observed peaks in H_3^+ density and emission occurring at a later time than the peak precipitation expected at 90° PPO phase.

38 1 Introduction

39 1.1 Ionosphere

40 Saturn's ionosphere is thought to be dominated by the positive ions H^+ and H_3^+ between 900
41 - 3000 km altitude and by hydrocarbon ions (e.g. C_3H_5^+) between 500 - 900 km altitude, along
42 with their companion electrons, which maintain the ionosphere's quasi-neutrality (*Moses and*
43 *Bass*, 2000). Co-located with this is the thermosphere, the charge-neutral component of the
44 upper atmosphere, which is composed chiefly of H and H_2 . Charged particles in the ionosphere
45 are continuously generated by ionising the otherwise neutral thermosphere through two main
46 mechanisms. The first, photo-ionisation by solar extreme ultra-violet (EUV) radiation, acts
47 across the entire sunlit portion of the planet (the dayside). The second, electron impact ioni-
48 sation, acts primarily in the polar regions of the planet. Both mechanisms also electronically,
49 vibrationally and rotationally excite the atmospheric constituents, which in turn de-excite
50 and emit photons. The emissions from these mechanisms are 'auroral' emissions and occur at
51 multiple wavelengths including infrared (IR), visible and ultraviolet (UV). This paper focuses
52 primarily on the infrared emissions emanating from the molecular ion H_3^+ near the poles of the
53 planet.

54 Saturn's ionosphere lies at the base of the planetary magnetosphere, a region formed by
55 the confinement of the planetary magnetic field by the solar wind. Closed field lines extend
56 in the equatorial region to distances $\sim 22 R_S$ (R_S is Saturn's 1 bar equatorial radius, equal
57 to 60,268 km) on the dayside (*Radioti et al.*, 2013), while open field lines stretch into a long
58 magnetic tail downstream from the planet on the nightside. From estimates of the open flux
59 in the magnetotail, the boundary between open and closed field lines in the ionosphere typi-
60 cally lies at around planetocentric co-latitude $\sim 15^\circ$ in each hemisphere (*Badman et al.*, 2006),
61 the difference between the two reflecting the north-south quadrupole asymmetry of Saturn's
62 planetary magnetic field (*Burton et al.*, 2010). In general it is expected that field-aligned cur-

63 rents flow down into the ionosphere over the polar field region due to the sub-corotation of
 64 plasma on open field lines and in the outer magnetosphere (*Bunce et al.*, 2008). The current
 65 then flows from the pole towards the equator in both hemispheres as ionospheric Pedersen
 66 currents, before returning up the field lines to the magnetosphere at lower latitudes as the
 67 flow returns to near-rigid corotation with the planet (e.g. *Cowley and Bunce*, 2003; *Cowley*
 68 *et al.*, 2004). The main auroral oval emissions are related to the latter ring of upward current
 69 (downward electron precipitation). The auroral oval is thus expected to lie in the region just
 70 equatorward of the open-closed boundary where the plasma angular velocity rises from low
 71 values on open lines towards rigid corotation on closed lines. The main oval is in general taken
 72 to correspond to the region between co-latitudes of $\sim 10^\circ$ and $\sim 20^\circ$ in both hemispheres (see,
 73 e.g., *Carbary*, 2012, and references therein). Auroral emissions are also sometimes observed
 74 in the poleward region, likely associated with solar wind-magnetosphere coupling dynamics
 75 at the magnetopause boundary of the magnetosphere (e.g. *Meredith et al.*, 2014). Here we
 76 present new observations of H_3^+ obtained with the Keck telescope in April 2013 using similar
 77 methodology to that employed by *O'Donoghue et al.* (2014). These observations measure the
 78 northern and southern main auroral ovals simultaneously as in the previous study, but this
 79 time they take place over three days instead of one, allowing for a wider ranging analysis
 80 of short term auroral behavior. In addition, due to the developing northern spring season at
 81 Saturn, the dataset presented here also includes and discusses simultaneous measurements of
 82 both the northern polar aurora as well as the midnight main auroral oval, owing to the viewing
 83 geometry at the time of the observations.

84 1.2 The H_3^+ probe at Saturn

85 The molecular ion H_3^+ is produced by the reaction $\text{H}_2 + \text{H}_2^+ \longrightarrow \text{H}_3^+ + \text{H}$ (*Oka*, 2006). The
 86 reaction time (the ion chemistry timescale) varies from 10 seconds at 800 km altitude to 1000

87 seconds for altitudes near 2000 km (*Badman et al.*, 2014). The lifetime of H_3^+ is proportional
88 to its temperature, inversely proportional to the ionospheric electron density and has been
89 previously quoted as 500 seconds (*Melin et al.*, 2011). During this lifetime, H_3^+ becomes ther-
90 mally excited to a higher rotational-vibrational (ro-vibrational) state by neighboring molecules
91 on timescales of 10^{-2} s, which is approximately the same time for the ion to relax to a lower
92 state and emit a photon. The discrete emission line spectra of H_3^+ make it a useful probe of
93 the conditions in Saturn’s ionosphere for two reasons. The first is that H_3^+ parameters such as
94 column-integrated temperature, density and power output (hereafter, total emission) can be
95 derived from it (e.g. *Miller et al.*, 2006; *Melin et al.*, 2014). Secondly, it is considered to be in
96 local thermodynamic equilibrium (LTE) - or at least quasi-LTE - with its surroundings (*Miller*
97 *et al.*, 1990; *Moore et al.*, 2008), meaning that the ion temperature is equivalent to the neutral
98 temperature.

99 Using the ground-based 3.8-metre United Kingdom InfraRed Telescope (UKIRT), the south-
100 ern auroral H_3^+ temperature was found to be 380 ± 70 K in 1999 and 420 ± 70 K in 2004 by *Melin*
101 *et al.* (2007). Later, in 2007, the Visual and Infrared Mapping Spectrometer (VIMS) (*Brown*
102 *et al.*, 2004) on board Cassini was used to derive a southern polar auroral H_3^+ temperature of
103 (on average) 590 ± 30 K over a period of 10 hours (*Stallard et al.*, 2012a). Measurements of the
104 southern auroral oval at equinox in 2009, also obtained by Cassini VIMS, yielded average tem-
105 peratures of ~ 410 K (*Lamy et al.*, 2013). The first conjugate northern and southern main oval
106 H_3^+ temperatures were measured at high spatial resolution in 2011 using the 10-metre W.M.
107 Keck II (hereafter, Keck) telescope by *O’Donoghue et al.* (2014). The 10 spectral images, when
108 co-added, yielded an average main auroral H_3^+ temperature of 583 ± 13 K (south) and 527 ± 18
109 K (north) over a ~ 2 hour period. Throughout this time interval the spectra gave temperatures
110 that varied by tens of Kelvins; this was a similar variability to the uncertainties, so it may
111 be considered real or due to noise. In the neutral thermosphere near the exobase (~ 1900 km
112 altitude above the 1 bar surface), solar occultations were performed using the Cassini ultra-

113 violet imaging spectrometer (UVIS) to derive temperatures (*Koskinen et al.*, 2013), yielding
114 temperatures of 370 K to 540 K from low- to high(auroral)-latitudes, respectively. The inter-
115 hemispheric temperature asymmetry measured by *O'Donoghue et al.* (2014) was postulated to
116 be the result of an inversely proportional relationship between magnetic field strength and the
117 total heating rate. Due to the lower magnetic field strength in the south, the area undergoing
118 heating is larger in the south than in the north (see *O'Donoghue et al.*, 2014, for a more detailed
119 discussion). Whilst the thermospheric temperatures at high latitudes can mostly be explained
120 via auroral region Joule heating (*Cowley et al.*, 2004), the low-latitude high temperatures re-
121 main difficult to explain theoretically. For example, exospheric temperatures are modeled to
122 be 143 Kelvin on the basis of solar EUV heating alone, yet observations show the exosphere
123 to be ~ 400 K (at sub-auroral latitudes) (*Yelle and Miller*, 2004; *Koskinen et al.*, 2013). *Smith*
124 *et al.* (2007) and *Mueller-Wodarg et al.* (2012) have explored the idea that heat is meridionally
125 transported down from the poles to the equator, but conclude that auroral heating actually
126 provides a net cooling effect at low latitudes. This is caused by a circulation pattern in which
127 high altitude heating (by ion drag) causes equatorward flows. The flow is balanced by the
128 continuity equation at low altitudes in the form of poleward flows, which themselves require
129 there be an upwelling of material from below. It is this upwelling material that expands and
130 cools adiabatically, leading to the counter intuitive effect of low latitude cooling, despite there
131 being a nearby heating source (*Smith et al.*, 2007). Thus, at present, it appears some addi-
132 tional source of energy is required to explain equatorial temperatures. One suggestion is the
133 breaking of gravity waves in the thermosphere, but this is modeled to account for temperature
134 enhancements of (at most) ~ 10 's of K (*Barrow and Matcheva*, 2013). A final source of note is
135 the low-latitude precipitation along the magnetic field lines conjugate to the rings known as
136 'ring rain'; it is possible that this is also associated with a low-latitude current system between
137 the rings and the planet, but as yet such currents have not been directly observed (*O'Donoghue*
138 *et al.*, 2013).

140 In 1980 both Voyager 1 and 2 spacecraft measured bursts of nonthermal radio emission which
 141 emanated from Saturn - specifically they are likely from the northern hemisphere: the period
 142 of these bursts were ~ 10.67 hours and taken (provisionally) to be the intrinsic rotation period
 143 of the planet (*Kaiser et al.*, 1980). However, more recently, during Saturn's pre-equinoctial
 144 southern summer between 2004 - 2008, the Cassini spacecraft has measured Saturn kilometric
 145 radiation (SKR) from both the northern and southern hemispheres, finding them to exhibit
 146 different periods: ~ 10.6 hours in the north and ~ 10.8 hours in the south (although these rates
 147 are still changing over time) (*Gurnett et al.*, 2009). These emissions, together with magnetic
 148 field perturbations observed within the magnetosphere, are inferred to be associated with
 149 two independent current systems rotating in the northern and southern hemispheres with
 150 slightly differing periods that vary slowly with Saturn's seasons (see, e.g. *Andrews et al.*,
 151 2008, 2010; *Southwood*, 2011; *Provan et al.*, 2009, 2012, and references therein). Following the
 152 recent discussion by *Southwood and Cowley* (2014), the empirically-determined current system
 153 associated with the northern ionosphere, of primary interest here, is shown in Figure 1, in
 154 a view looking down on the northern pole (a similar current system also flows in the south)
 155 (*Hunt et al.*, 2014). In this diagram the solid lines and symbols show the currents, while the
 156 dotted lines represent the associated magnetic field perturbations above the Pedersen layer of
 157 the ionosphere required by Ampère's law. The primary current system consists of field-aligned
 158 currents that flow down into the ionosphere on the right of the diagram (circled crosses on the
 159 inner dashed line ring), across the polar ionosphere as Pedersen currents directed from right
 160 to left, and out of the ionosphere as field-aligned currents on the left of the diagram (circled
 161 dots on the black dashed line ring). Secondary field-aligned currents of lesser magnitude and
 162 opposite polarity also flow on the outer ring, which serve to limit the field perturbations to
 163 the interior region. This current system then rotates with the northern period, ~ 10.64 h at

164 the time of our observations (compared with ~ 10.69 h for the southern SKR period). Position
 165 with respect to the rotating pattern is defined by the northern PPO phase function Ψ_N , which
 166 increases clockwise around the diagram in Figure 1. Enhanced upward currents, associated
 167 with enhanced electron precipitation and auroral emissions, are expected to occur for $\Psi_N \approx$
 168 90° (modulo 360°), while enhanced downward currents, likely associated with suppression of
 169 precipitation and emissions, are expected for $\Psi_N \approx 270^\circ$.

170 Empirically, the orientation of the system at any time is determined through examination of
 171 the related magnetic field oscillations. In particular, if we consider the magnetic perturbations
 172 between the two current rings (dotted lines in Figure 1), mapped along quasi-dipolar field lines
 173 into the equatorial magnetosphere, it will be seen that these transform into a quasi-uniform
 174 field in which the perturbation field points radially outward from the planet at $\Psi_N \approx 0^\circ$, radially
 175 inward at $\Psi_N \approx 180^\circ$, and has positive and negative azimuthal components (with respect to
 176 the northern spin/magnetic pole) at $\Psi_N \approx 90^\circ$ and 270° , respectively. Magnetic oscillations
 177 observed in the equatorial magnetosphere are then analysed to determine the azimuth with
 178 respect to noon at which the northern quasi-uniform perturbation field points radially outward
 179 at any instant of time, $\Phi_N(t)$, thus also defining the azimuth where the northern PPO phase
 180 Ψ_N takes the value zero (modulo 360°) at that time. The northern PPO phase as a function
 181 of azimuth and time is thus given by

$$182 \quad \Psi_N(\phi, t) = \Phi_N(t) - \phi, \tag{1}$$

183 where ϕ is the azimuth in degrees with respect to noon of any observation point (equiva-
 184 lent to local time), and $\Phi_N(t)$ is determined empirically, with rotation period given by τ_N
 185 $= 360^\circ / (d\Phi_N/dt)$ and with Φ_N expressed in degrees. The function $\Phi_N(t)$ employed here is
 186 that determined from Cassini magnetic field data by *Provan et al.* (2014). Signatures of this
 187 planetary period oscillation from the auroral region were first noted from the Voyager 1 and
 188 2 spacecraft's UV photometer data by *Sandel and Broadfoot* [1981] and *Sandel et al.* [1982].

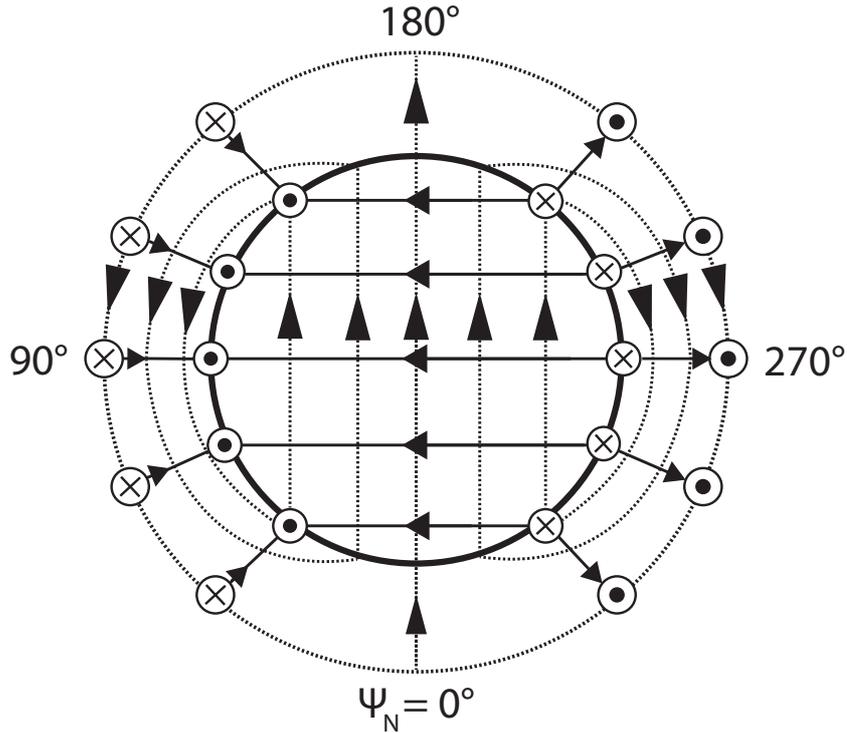


Fig. 1. Sketch of the form of the currents (solid lines and symbols) and perturbation magnetic fields (dotted lines) associated with the northern system PPOs, in a view looking down on Saturn’s northern ionosphere from above. The principal field-aligned currents flow across the inner ring: into the ionosphere on the right (circled crosses), and out of the ionosphere on the left (circled dots), joined by ionospheric Pedersen currents flowing from right to left across the polar ionosphere. Secondary field-aligned currents of smaller magnitude and opposite polarity flow on the outer boundary of the current system, confining the perturbation field to the interior region. The current system rotates anti-clockwise with the northern PPO period τ_N . Azimuth with respect to the current system is defined by the phase function $\Psi_N(\phi, t)$ as shown in the figure (equation (1)), increasing clockwise around the diagram.

189 Using Cassini VIMS, *Badman et al.* [2012a] discovered that the H_3^+ auroral intensity follows a
 190 sinusoidal function with PPO phase, with H_3^+ peak intensity occurring in the north between
 191 $\Psi_N \approx 0 - 45^\circ$, before the expected maximum peak intensity associated with enhanced electron
 192 precipitation at $\Psi_N \approx 90^\circ$.

2 Observations

The observations presented here used the 10-m Keck telescope situated on Mauna Kea, Hawaii. They were designed to be an integral part of the Saturn Auroral Observing Campaign of 2013 (this *Icarus* special issue), such that they overlap observations performed by the Cassini spacecraft, Hubble Space Telescope, and the NASA Infrared Telescope Facility (IRTF). The observations took place on the 19, 20, and 21 April and are summarized in Table 1. In this table the quoted times are the actual observing time on Earth (i.e. not corrected for light-travel time from Saturn to Earth) and the ‘seeing’ column refers to blurring of the received light by the Earth’s atmosphere. The quoted central meridian longitudes (CMLs) are from the Saturn system III longitude system [*Kaiser et al.* 1980]. Emissions from these CMLs are light travel time corrected, i.e. the ~ 73 minutes time delay has been accounted for in the results here. During these dates, Saturn was at opposition with respect to the Earth-Sun line with its northern hemisphere tilted towards the Earth and the Sun with both a sub-Earth and sub-solar latitude (coincidentally) of 18.3° , i.e. in conditions of Saturn’s northern spring (summer solstice occurs in 2017). In the previous work, Saturn had a sub-Earth latitude of 8.2° [*O’Donoghue et al.* 2014].

Date	Start UT	End UT	Saturn integration*	CML range	Seeing
19 April	10:55:00	13:11:50	40 min (8)*	43 - 120°	0.4''
20 April	12:18:42	13:18:39	20 min (4)*	181 - 215°	0.45''
21 April	10:40:05	13:24:41	55 min (11)*	217 - 309°	0.6''

Table 1

Summary of Keck telescope observations in April 2013. *Total time spent observing Saturn itself; the number in parentheses is the number of 5-minute co-additions used.

The instrument used on the Keck telescope was the near infrared spectrometer (NIRSPEC) [*McLean et al.* 1998], which has a spectral resolving power of $R = \lambda/\Delta\lambda \sim 25,000$ and thus

211 provides a minimum resolution of (e.g.) $\Delta\lambda \approx 1.59 \times 10^{-4} \mu\text{m}$ at $3.975 \mu\text{m}$. The wavelength
 212 range used here is between 3.95 and $4.0 \mu\text{m}$, which covers the Q-Branch ($\Delta J=0$) ro-vibrational
 213 transition lines of H_3^+ . Saturn's axis of rotation is measured to be co-aligned with the magnetic
 214 axis to within $\sim 0.1^\circ$ uncertainty [*Burton et al.* 2010]. Taking advantage of this symmetry, the
 215 spectrometer slit was orientated in a north-south direction on Saturn as shown in Figure 2.
 216 The planet is then allowed to rotate beneath the slit whilst spectral images are taken along
 217 the noon-midnight meridian plane. The slit measures $0.432''$ width by $24''$ length with a pixel
 218 on the CCD corresponding to $0.144''$ squared on the sky, as in Figure 2.

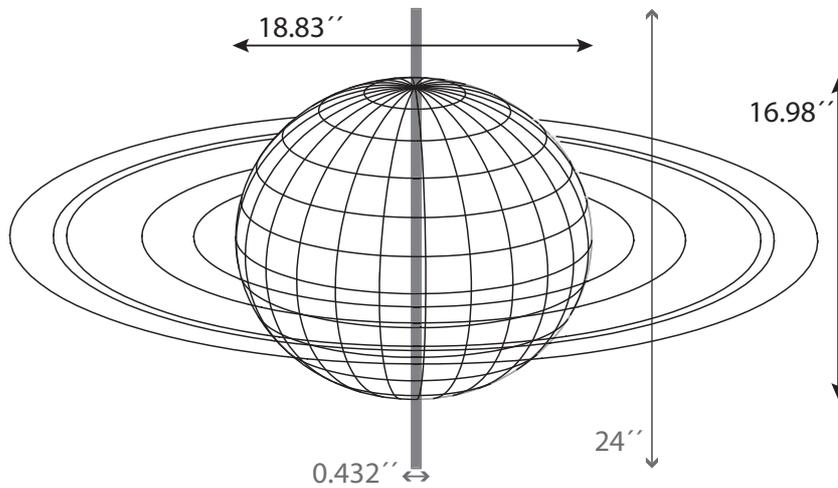


Fig. 2. **Saturn as observed with Keck, April 21 2013.** Saturn's sub-Earth latitude was 18.3° during the observations. The arrowed lines show the angular extent of Saturn and the dimensions of the NIRSPEC spectral slit in seconds of arc.

219 Owing to this viewing geometry we are afforded the ability to collect data from four distinct
 220 latitudinal ranges:

- 221 (1) Northern midnight main oval (NMMO): $8 - 15^\circ$ co-latitude (Nightside)
- 222 (2) Northern polar cap (NPC): $0 - 6^\circ$ co-latitude (Day - Nightside)
- 223 (3) Northern main oval (NMO): $8 - 22^\circ$ co-latitude (Dayside)
- 224 (4) Southern main oval (SMO): $18 - 22^\circ$ co-latitude (Dayside)

225 where dayside and nightside correspond to regions sunward and anti-sunward of the krono-
226 graphic north pole, respectively. These regions of interest are shown in Figure 3 and note that
227 they all remain lit by the Sun. They were selected (as close as the viewing geometry allowed) to
228 coincide with the approximate statistical locations of the northern and southern main auroral
229 ovals between $\sim 10 - 25^\circ$, and the polar cap between $\pm 10^\circ$ of the north pole [*Badman et al.*
230 2006; *Carbary* 2012]. These regions are associated with internal and external forcing on the
231 Saturnian magnetosphere, respectively, as discussed in the introduction. An example of the
232 viewing geometry limitation is at the NPC - here, the spatial resolution of one pixel on the
233 detector corresponds to $\sim 3^\circ$ latitude. In addition, and applicable to the whole spectral image,
234 atmospheric seeing will smear the signal received across multiple pixels. Although the amount
235 of pixels smeared is constant within the image, the range of latitudes represented by a given
236 pixel diminishes with increasing latitude. This cross-contamination by light from neighbouring
237 pixels is taken into account by creating a small separation of between $\sim 0.144 - 0.288$ seconds
238 of arc (1 - 2 pixels) between the different regions listed above.

239 Each individual spectral image consists of twelve 5-s integrations, creating exposures 60 s
240 long, which are of both Saturn ‘A’ and sky ‘B’ frames, with the telescope slewing between each
241 in the sky in an ABBA pattern. Standard astronomical reduction techniques are employed
242 to clean the observed spectral images, which include an A-minus-B subtraction in which the
243 Earth’s sky emissions are removed from the Saturn spectra, and a star flux calibration. The flux
244 calibration measures the spectrum of a black body emitting star (A0) in order to account for
245 the wavelength dependent absorption of light by the Earth’s atmosphere, whilst also converting
246 the CCD photon count into physical photon flux. The star used in this work was HR 5717.
247 Other reduction procedures include a dark current subtraction and dividing by a ‘flat field’.
248 Together, these account for thermal emissions at the detector and defects on the CCD chip and
249 optics, respectively. The reduced spectral images are then co-added into groups of five spectra
250 (see Table 1) in order to create a single higher signal to noise (S/N) ratio image. However,

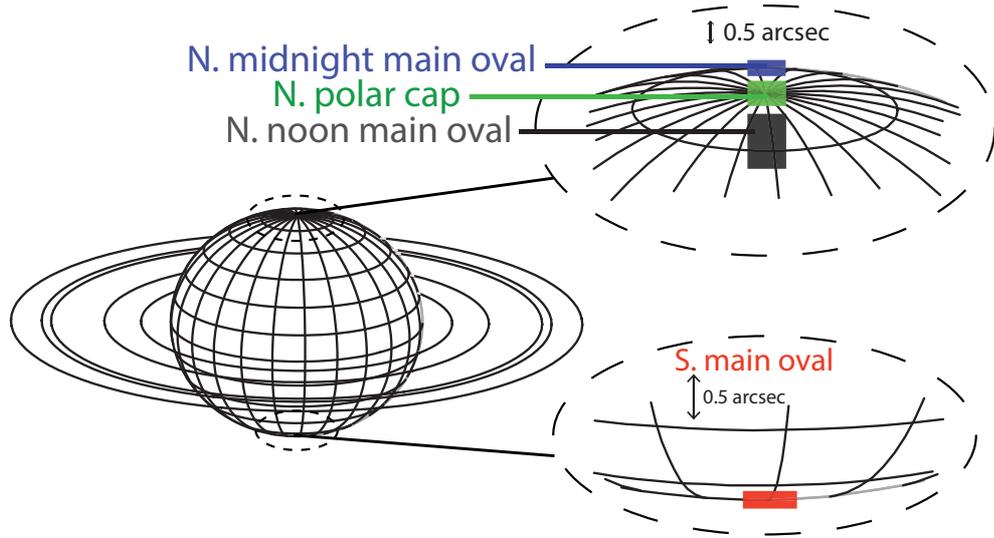


Fig. 3. **Regions of interest on Saturn.** Four distinct color-coded areas are illustrated, corresponding to the regions listed in the text. The chosen color scheme will be used in subsequent figures for clarity. Note that the different colored blocks (not to exact scale) are separated slightly in the north, to avoid cross-contamination introduced by the effects of telluric seeing. Longitude and latitude grid lines represent 15 degree spacings.

251 these spectra are obtained at different times typically within a $\sim 15 - 20$ minute range; this
 252 is chiefly because the A frames are often separated by B frames, but more general observing
 253 time overheads cause this time window to vary, e.g. the telescope slewing time between the A
 254 and B frame positions, losses in tracking or human error. Within these time ranges we thus
 255 typically obtain a swath of data spanning $8 - 11^\circ$ in longitude as the planet rotates beneath
 256 the slit. In this work we assume an optically thin atmosphere in and above the ionosphere;
 257 this assumption was used and tested by *Lam et al.* [1997] to be valid.

258 An example of a reduced spectral image (x-axis wavelength, y-axis spatial dimension) which
 259 has been co-added from all 5-minute integrations on April 21 is shown in Figure 4. In this figure
 260 there are three main sources of radiation highlighted: the reflection of sunlight from the lower
 261 atmosphere, the continuum reflection of sunlight from the rings, and discrete H_3^+ emission
 262 lines. The ability to measure H_3^+ emissions is aided by the fact that hydrocarbons also absorb

263 sunlight at different wavelengths; these are the dark regions on the body of the planet between
 5 and 19 arcseconds in Figure 4.

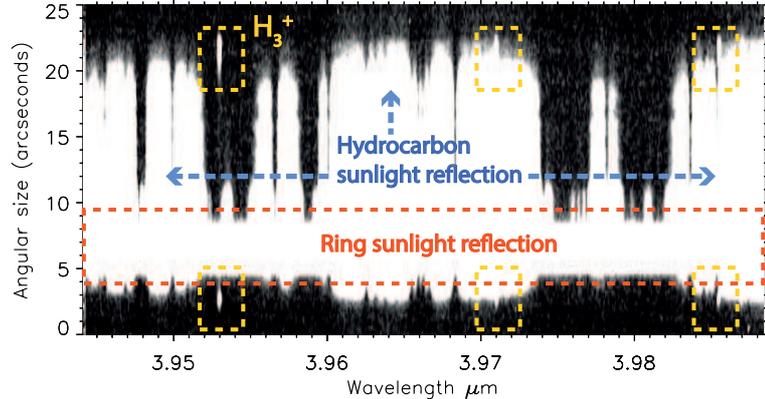


Fig. 4. **An image of the spectrum of Saturn taken at local Saturn noon.** This is the co-addition of all eleven 5-minute integrations on 21 April 2013. The wavelength range is shown on the horizontal axis and the angular size in the sky is shown on the vertical axis. North is at the top of the image and south is at the bottom. Discrete H_3^+ emission line spectra are inside the yellow dashed boxes in the form of white vertical lines (white being high light intensity, black being low/none). From left to right these lines are the $Q(1,0^-)$, $Q(2,1)$ and $Q(3,0)$ lines described in the main text. Hydrocarbons such as methane absorb solar radiation (creating the black background) between the auroral regions. The white bar of emission centered at $\sim 6''$ is the continuum reflection of sunlight from the rings. The remaining white pixels are due to sunlight reflected by hydrocarbons and other molecules.

264

265 3 Data analysis

266 For a given temperature, a discrete H_3^+ emission line will emit at a given intensity. We
 267 produce a theoretical spectrum of multiple lines from a line list of H_3^+ emission for thousands
 268 of different temperatures [see e.g. *Neale et al.* 1996; *Melin et al.* 2014]. The relative intensities of
 269 multiple discrete H_3^+ emission lines (i.e. a set of line ratios) represent the effective temperature
 270 of H_3^+ in quasi-LTE. An example of an observed spectral profile is shown in Figure 5 by the
 271 black crosses. Three Q-branch ($\Delta J = 0$) intensity peaks are visible from left to right in this

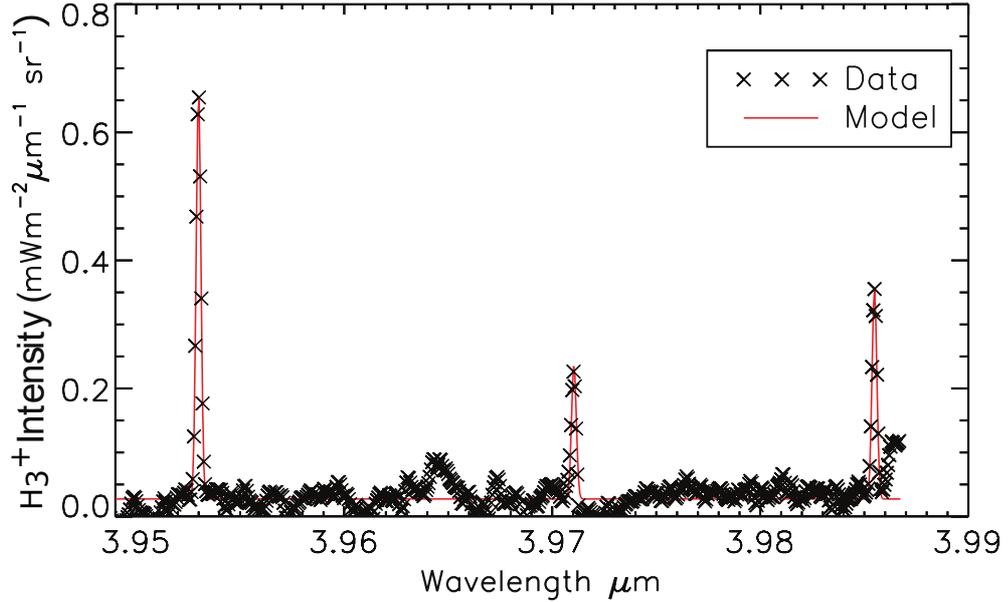


Fig. 5. **Model fit to H_3^+ intensity as a function of wavelength.** This spectral profile is produced from the co-add of all northern main oval images on April 21 (NMO; 8 - 22° co-latitude). The x- and y-axes show wavelength and intensity of H_3^+ emission, respectively. The latter is indicated by the black crosses for the observed emission and the model fit to the spectrum is shown in red. The temperature derived for this spectral profile is 404 ± 11 K.

272 figure; Q(1,0⁻), Q(2,1) and Q(3,0). These have transition energies ω between upper (j) and
 273 lower (i) ro-vibrational energy states of $\omega_{i,j} = 2529.721 \text{ cm}^{-1}$, $\omega_{i,j} = 2514.619 \text{ cm}^{-1}$ and $\omega_{i,j} =$
 274 2509.074 cm^{-1} , respectively (further transition line information is available in Table 1 of *Kao*
 275 *et al.* [1991]). The modeled theoretical spectrum is reproduced for a variety of temperatures
 276 until a close match is found to the observed spectrum by least-squares fitting. In other words,
 277 the effective column-integrated temperature of H_3^+ is found by comparing the observed line
 278 ratios to model values.

279 Emission by H_3^+ depends upon its temperature with, in general, a higher temperature lead-
 280 ing to a spectral transition line of higher intensity. The emission we observe at the detector
 281 (following the data reduction) is representative of a line-of-sight column-integrated quantity
 282 of molecules: the column density. Thus by dividing the observed intensity by the intensity of a

283 single molecule we can determine the number of molecules in the column, in units of molecules
 284 per square metre (m^{-2}).

285 The effective total emission of H_3^+ is the result of the multiplication of the integrated emission
 286 per molecule across all wavelengths by the column density, giving a measurement of the total
 287 emitted power by H_3^+ as follows [*Lam et al.* 1997]:

$$E(\text{H}_3^+) = E_{mol}(\text{H}_3^+, T) \times N(\text{H}_3^+, T) , \quad (2)$$

288 where $E_{mol}(\text{H}_3^+, T)$ is the integrated intensity of a H_3^+ line between 0.75 - 22 microns

$$289 \quad E_{mol}(\text{H}_3^+) = a + bT + cT^2 + dT^3 + eT^4 . \quad (3)$$

290 modeled for a particular temperature, T . Parameters a to e are partition function constants
 291 detailed and given by *Miller et al.* [2010]. As such, it is a direct measure for the rate of cooling
 292 of the ionosphere/thermosphere by H_3^+ , which is itself responsible for some of the cooling in
 293 the thermosphere [*Grodent et al.* 2001; *Raynaud et al.* 2004; *Miller et al.* 2010].

294 The line-of-sight column density attains a useful physical meaning if it is corrected to the
 295 altitude of a column that extends vertically from Saturn's surface. Thus each such measurement
 296 needs to be reduced by some factor dependent upon the angle to the local vertical of the
 297 observation. Observations by the Cassini spacecraft show that the majority of the H_3^+ intensity
 298 is located within the 800 - 1400 km range of altitudes above the 1 bar pressure surface [*Stallard*
 299 *et al.* 2012b]. Models are in agreement with this, predicting that the majority of ionospheric
 300 H_3^+ ions (approximately >90% by number density) are located in this same 600 km range of
 301 altitudes [e.g. *Mueller-Wodarg et al.* 2012]. By considering two oblate spheroids (with elliptical
 302 cross-sections) of Saturn tilted at 18.3° relative to the observer, the inner spheroid being the
 303 1-bar pressure surface of Saturn plus 800 km, and the outer spheroid being at plus 1400 km
 304 altitude of the same surface, we calculated the depth of the column we observe as a function

305 of latitude. The observed line of sight column of atmosphere becomes larger nearer the poles,
306 compared to at the equator, and this is corrected for by reducing the measured intensity as a
307 function of latitude to a normalised value.

308 The spectrum of H_3^+ can be described as a ‘spectral function’: this function is a sum of
309 Gaussian fits to all of the ro-vibrational transition lines and depends on the temperature
310 and density of H_3^+ . The temperature and density uncertainties from this are found by applying
311 Cramer’s rule, whilst the uncertainty in total emission is found by using basic error propagation
312 formulae [see *Melin* 2006; *Melin et al.* 2014, and references therein]. As the temperature and
313 density parameters are found using a least-squares fit embedded within the H_3^+ fitting routine,
314 they are an indicator of the quality of the spectral fit.

315 The most optimal (lowest) seeing achieved herein is $0.4''$, which amounts to 2560 km perpen-
316 dicular to the line-of-sight at a distance of 8.826 astronomical units (the Earth-Saturn distance
317 on April 20). Therefore, even in the extreme case of a measurement of H_3^+ on or near the Sat-
318 urnian limb, we are still capturing the entire column of altitude in which H_3^+ is distributed
319 above the 1-bar surface. This means that variability in H_3^+ parameters that we report herein
320 should be considered due to variations in latitude and longitude and not in altitude which is
321 column-integrated at any location.

322 4 Results and discussion

323 The total time spent observing in this campaign, including sky exposures of the sky, cali-
324 brations and general time-overheads (e.g. moving the telescope), was 361 minutes. The total
325 integration time on Saturn itself was 115 minutes. The exposure times and values of tempera-
326 ture, column density, and total emission for the northern and southern main ovals (NMO and
327 SMO) and northern polar cap and midnight oval (NPC and NMMO) are shown in Tables 2,

328 3 and 4 for April 19, 20 and 21, respectively. The start and end universal times (UT) in the
329 tables correspond to the start of the first Saturn exposure (A frame) and the end of the final
330 (fifth) A frame; as mentioned earlier, observing time overheads do not permit a continuous
331 5-minute acquisition.

332 Average parameters for each day are shown in two different ways in this section. The
333 first is the average of all individually model-fitted spectra for a given parameter over a given
334 observation night; these are represented by the dashed horizontal lines in each of the figures (6
335 - 9 inclusive). Note that the values in the first row of Table 2 have unusually high uncertainties,
336 perhaps due to passing cirrus clouds during the observations; as such, they are not used when
337 calculating the averages. A second type of averaging, the ‘co-average’, is found by fitting a
338 model H_3^+ spectrum to the co-addition of all spectral images from each region for each day.
339 This ensures that the maximum possible S/N is obtained prior to fitting itself. These co-added
340 averages have higher S/N and lower uncertainty than an individually fitted spectral image and
341 are given in Tables 2, 3 and 4, although the difference between the two types of averaging are
342 small.

Start (UT)	End (UT)	T_{NMO} (K)	T_{NPC} (K)	T_{NMMO} (K)	T_{SMO} (K)	CD_{NMO} (10^{15} m^{-2})	CD_{NPC} (10^{15} m^{-2})	CD_{NMMO} (10^{15} m^{-2})	CD_{SMO} (10^{15} m^{-2})	E_{NMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{NPC} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{NMMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{SMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)
10:55	11:18	994 ± 900	925 ± 587	699 ± 212	453 ± 67	0.04 ± 0.06	0.03 ± 0.05	0.03 ± 0.05	0.5 ± 0.5	0.70 ± 0.09	0.37 ± 0.09	0.12 ± 0.02	0.11 ± 0.07
11:23	11:50	389 ± 34	356 ± 32	501 ± 48	466 ± 39	7.4 ± 5.3	10.2 ± 8.6	0.3 ± 0.2	0.7 ± 0.4	0.40 ± 0.37	0.23 ± 0.46	0.14 ± 0.02	0.18 ± 0.03
11:51	12:05	449 ± 38	394 ± 37	506 ± 58	361 ± 30	2.5 ± 1.5	4.1 ± 3.2	0.3 ± 0.2	6.3 ± 4.9	0.46 ± 0.14	0.25 ± 0.21	0.12 ± 0.02	0.17 ± 0.20
12:09	12:22	376 ± 32	496 ± 41	438 ± 50	583 ± 54	10.3 ± 7.8	1.0 ± 0.5	0.7 ± 0.6	0.2 ± 0.1	0.41 ± 0.47	0.39 ± 0.05	0.11 ± 0.05	0.24 ± 0.01
12:23	12:36	396 ± 26	428 ± 32	435 ± 38	498 ± 46	9.0 ± 4.9	2.8 ± 1.6	0.9 ± 0.6	0.5 ± 0.3	0.56 ± 0.25	0.34 ± 0.10	0.13 ± 0.04	0.20 ± 0.03
12:41	12:53	417 ± 32	372 ± 32	471 ± 54	398 ± 51	5.2 ± 3.2	7.8 ± 6.0	0.6 ± 0.5	1.9 ± 2.1	0.52 ± 0.21	0.27 ± 0.31	0.16 ± 0.04	0.13 ± 0.14
12:54	13:07	382 ± 49	451 ± 33	395 ± 34	479 ± 45	11.4 ± 11.4	2.0 ± 1.0	2.1 ± 1.5	0.6 ± 0.4	0.52 ± 0.62	0.38 ± 0.07	0.13 ± 0.08	0.18 ± 0.03
13:11	13:24	407 ± 22	467 ± 38	436 ± 45	444 ± 41	7.9 ± 3.5	1.4 ± 0.8	0.7 ± 0.6	1.0 ± 0.7	0.64 ± 0.17	0.37 ± 0.06	0.10 ± 0.05	0.17 ± 0.05
Co-average*:	402 ± 20	441 ± 16	466 ± 20	442 ± 23	6.4 ± 2.5	2.0 ± 0.5	0.5 ± 0.1	0.9 ± 0.4	0.47 ± 0.02	0.31 ± 0.04	0.12 ± 0.02	0.15 ± 0.03	0.15 ± 0.03
Mean value	404 ± 13	423 ± 13	455 ± 18	460 ± 17	7.7 ± 2.3	4.2 ± 1.6	0.8 ± 0.3	1.5 ± 0.7	0.50 ± 0.12	0.32 ± 0.09	0.13 ± 0.02	0.17 ± 0.03	0.17 ± 0.03

Table 2

Saturn's auroral/polar properties as a function of time on 19 April 2013. All uncertainties shown are one standard deviation (i.e. 1-sigma errors). T, CD and E are temperature, column density and total emission of H_3^+ , respectively. *Co-averages are co-add averages formed from applying a model fit to the co-addition of all spectra from the night, rather than of the individual values, whilst the mean values are drawn from the table. Note that the first row is not used in the latter average due to very high uncertainties.

Start (UT)	End (UT)	T_{NMO} (K)	T_{NPC} (K)	T_{NMMO} (K)	T_{SMO} (K)	CD_{NMO} (10^{15} m^{-2})	CD_{NPC} (10^{15} m^{-2})	CD_{NMMO} (10^{15} m^{-2})	CD_{SMO} (10^{15} m^{-2})	E_{NMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{NPC} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{NMMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{SMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)
12:18	12:31	461 ± 42	426 ± 35	476 ± 33	475 ± 41	2.3 ± 1.5	2.5 ± 1.6	0.7 ± 0.3	0.7 ± 0.4	0.54 ± 0.13	0.31 ± 0.11	0.2 ± 0.03	0.21 ± 0.03
12:31	12:46	476 ± 37	423 ± 33	459 ± 32	460 ± 32	2.1 ± 1.1	2.8 ± 1.7	0.7 ± 0.4	1.1 ± 0.5	0.61 ± 0.09	0.32 ± 0.11	0.15 ± 0.03	0.24 ± 0.03
12:51	13:03	442 ± 34	377 ± 27	439 ± 31	496 ± 35	3.3 ± 1.9	7.1 ± 4.4	1.0 ± 0.5	0.6 ± 0.3	0.53 ± 0.14	0.29 ± 0.23	0.15 ± 0.03	0.23 ± 0.02
13:07	13:18	441 ± 31	503 ± 37	562 ± 47	441 ± 39	3.1 ± 1.6	0.8 ± 0.4	0.2 ± 0.1	1.2 ± 0.8	0.49 ± 0.13	0.34 ± 0.04	0.19 ± 0.01	0.19 ± 0.05
Co-average*	441 ± 22	423 ± 15	471 ± 17	454 ± 22	3.3 ± 1.3	2.8 ± 0.7	0.6 ± 0.2	1.0 ± 0.4	0.54 ± 0.09	0.31 ± 0.05	0.17 ± 0.01	0.22 ± 0.02	0.22 ± 0.02
Mean value	453 ± 20	434 ± 20	487 ± 22	468 ± 19	2.8 ± 0.9	3.6 ± 1.6	0.6 ± 0.3	0.9 ± 0.3	0.54 ± 0.07	0.31 ± 0.09	0.16 ± 0.02	0.22 ± 0.02	0.22 ± 0.02

Table 3

As Table 2, but for data obtained on 20 April 2013.

Start (UT)	End (UT)	T_{NMO} (K)	T_{NPC} (K)	T_{NMMO} (K)	T_{SMO} (K)	CD_{NMO} (10^{15} m^{-2})	CD_{NPC} (10^{15} m^{-2})	CD_{NMMO} (10^{15} m^{-2})	CD_{SMO} (10^{15} m^{-2})	E_{NMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{NPC} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{NMMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)	E_{SMO} ($10^{-5} \text{ Wm}^{-2} \text{ sr}^{-1}$)
10:40	10:53	375 ± 18	397 ± 25	525 ± 40	375 ± 30	23.9 ± 10.0	9.6 ± 5.0	0.4 ± 0.2	4.5 ± 3.2	0.89 ± 0.27	0.61 ± 0.14	0.23 ± 0.01	0.17 ± 0.14
10:54	11:07	415 ± 17	387 ± 22	388 ± 34	380 ± 21	9.8 ± 3.2	10.0 ± 5.0	3.6 ± 2.7	6.0 ± 3.0	0.95 ± 0.11	0.51 ± 0.15	0.19 ± 0.09	0.25 ± 0.09
11:12	11:26	384 ± 17	421 ± 21	384 ± 29	398 ± 25	17.3 ± 6.8	5.5 ± 2.1	3.9 ± 2.5	4.3 ± 2.3	0.83 ± 0.21	0.60 ± 0.07	0.18 ± 0.09	0.29 ± 0.07
11:28	11:41	380 ± 17	448 ± 25	393 ± 26	510 ± 36	20.2 ± 7.8	3.4 ± 1.4	3.0 ± 1.7	0.8 ± 0.4	0.87 ± 0.23	0.62 ± 0.06	0.18 ± 0.06	0.39 ± 0.02
11:46	11:59	437 ± 25	506 ± 28	491 ± 43	425 ± 26	5.7 ± 2.4	1.4 ± 0.5	0.7 ± 0.4	2.6 ± 1.2	0.85 ± 0.11	0.67 ± 0.03	0.28 ± 0.02	0.31 ± 0.04
12:00	12:13	364 ± 27	500 ± 35	628 ± 64	495 ± 38	19.1 ± 12.8	1.4 ± 0.6	0.1 ± 0.1	0.6 ± 0.3	0.53 ± 0.53	0.59 ± 0.04	0.24 ± 0.01	0.26 ± 0.02
12:18	12:31	477 ± 35	460 ± 26	439 ± 38	415 ± 28	2.2 ± 1.1	2.7 ± 1.1	1.2 ± 0.8	2.8 ± 1.5	0.68 ± 0.09	0.60 ± 0.05	0.19 ± 0.04	0.27 ± 0.06
12:33	12:46	402 ± 27	407 ± 19	448 ± 32	453 ± 22	8.3 ± 4.5	7.6 ± 2.8	1.1 ± 0.6	2.0 ± 0.7	0.59 ± 0.23	0.62 ± 0.08	0.19 ± 0.03	0.40 ± 0.02
12:51	13:03	391 ± 32	394 ± 26	448 ± 30	452 ± 20	8.3 ± 5.8	9.1 ± 5.1	1.2 ± 0.6	2.0 ± 0.7	0.46 ± 0.35	0.54 ± 0.16	0.21 ± 0.03	0.41 ± 0.02
13:05	13:18	423 ± 29	485 ± 29	495 ± 37	436 ± 22	5.3 ± 2.9	1.9 ± 0.8	0.6 ± 0.3	2.6 ± 1.0	0.60 ± 0.17	0.64 ± 0.04	0.23 ± 0.02	0.37 ± 0.03
13:23	13:36	420 ± 32	505 ± 28	483 ± 46	512 ± 33	4.9 ± 2.9	1.5 ± 0.5	0.5 ± 0.3	0.8 ± 0.3	0.52 ± 0.19	0.69 ± 0.03	0.17 ± 0.03	0.38 ± 0.02
Co-average*	404 ± 11	436 ± 9	460 ± 11	436 ± 10	9.2 ± 2.1	4.0 ± 0.6	0.9 ± 0.2	2.1 ± 0.4	0.70 ± 0.09	0.60 ± 0.02	0.20 ± 0.01	0.31 ± 0.01	0.31 ± 0.01
Mean value	409 ± 8	451 ± 9	460 ± 12	441 ± 9	10.1 ± 1.9	4.5 ± 0.8	1.6 ± 0.4	2.6 ± 0.5	0.70 ± 0.08	0.61 ± 0.03	0.21 ± 0.02	0.32 ± 0.02	0.32 ± 0.02

Table 4

As Table 2, but for data obtained on 21 April 2013.

344 In Tables 2 - 4, the co-added average temperatures in the NMO are lower than in the SMO
 345 on each day. The individually derived H_3^+ temperatures for the spectral images are shown in
 346 Figure 6, together with dashed lines which indicate the average value of all of the data points
 347 (i.e. not the same averages as in Tables 2 - 4, but the differences between the two are very
 348 small). *O'Donoghue et al.* [2014] found that over a period of ~ 2 hours the southern main
 349 auroral oval was on average 56 K hotter than its northern counterpart. This was attributed
 350 to the north-south asymmetry in magnetic field strength which leads to an overall larger total
 351 heating rate in the south, with the caveat being that their dataset was small and considered a
 352 snapshot of events at that time (in April 2011). In this work we have three similar snapshots
 353 over consecutive days, each appearing to support to the previous result that the SMO is warmer
 354 than the NMO by 10's of K when measured simultaneously for each of the days.

355 A summary of the effective average H_3^+ temperatures observed to date is presented in Table
 356 5. The considerable year-to-year variability is difficult to attribute to seasonal or solar cycle
 357 effects, such that variability on shorter time scales of minutes, hours, and days should be
 358 considered. This is discussed in Subsection 4.3 where we outline a likely reason for the several
 359 10's of Kelvin variability seen in the NMO temperatures.

360 Tables 2 - 4 also show that column densities are higher in the northern main oval than
 361 the southern by on average a factor of ~ 3 , as shown in Figure 7, though these have large
 362 uncertainties associated with them. A possible reason for a higher northern column density is
 363 the additional solar illumination in the north compared with that incident at the south; this
 364 yields a higher ionisation rate of H_2 and therefore an increase in H_3^+ production. Such an effect
 365 has previously been observed and also demonstrated using the 1-D Saturn Thermosphere Iono-
 366 sphere Model (STIM) by *O'Donoghue et al.* [2014]. All but one pair of values is in agreement

Date	T_{SMO} (K)	T_{NMO} (K)	T_{NPC} (K)	T_{NMMO} (K)	Source
Sept. 1999	380 ± 70	-	-	-	NASA IRTF, <i>Melin et al.</i> [2007]
Feb. 2004	420 ± 70	-	-	-	NASA IRTF, <i>Melin et al.</i> [2007]
July 2007	590 ± 50	-	-	-	Cassini VIMS, <i>Stallard et al.</i> [2012a]
Jan. 2009	410 ± 85	-	-	-	Cassini VIMS, <i>Lamy et al.</i> [2013]
April 2011	583 ± 13	527 ± 18	-	-	Keck, <i>O'Donoghue et al.</i> [2014]
April 2013	444 ± 18	416 ± 18	433 ± 13	466 ± 16	Keck, This work

Table 5

The average temperatures of Saturn's auroral regions obtained between 1999 and the 2013.

367 with this trend; at ~ 12 UT on April 19 in panel (a) the southern column density is higher.
368 The densities vary by up to an order of magnitude from day-to-day, with the major deviations
369 outside the ranges of uncertainty seen in panel (c).

370 The variability in column density is likely to be associated with changes in the energy flux
371 that is incident on the ionosphere, e.g. increased particle precipitation provides more ionization
372 and thus more H_3^+ . Similar variability in the energy flux has been attributed to variations in
373 H_3^+ aurora before using Cassini VIMS data [*Badman et al.* 2012b;a], and in patches of intense
374 UV emissions from H and H_2 [*Nichols et al.* 2009; *Grodent et al.* 2011; *Meredith et al.* 2013]. An
375 influx of particles at local noon may be the result of dayside reconnection events which occur
376 when the interplanetary magnetic field (IMF) is orientated northward, leading to the opening
377 of closed planetary magnetic field lines to the solar wind, causing a planetward influx of solar
378 particles [*Radioti et al.* 2011; 2013; *Badman et al.* 2013; *Meredith et al.* 2014; *Belenkaya et al.*
379 2014]. Alternatively, new parts of the main auroral oval, differing in their levels of activity, may
380 be rotating into view on the spectrograph slit. No correlations are found between the northern
381 and southern main ovals, despite sharing common (closed) magnetic field lines, and this is
382 consistent with recent Hubble Space Telescope (HST) observations which showed patches of
383 UV emission in the auroral main oval are present in one hemisphere, but absent from the

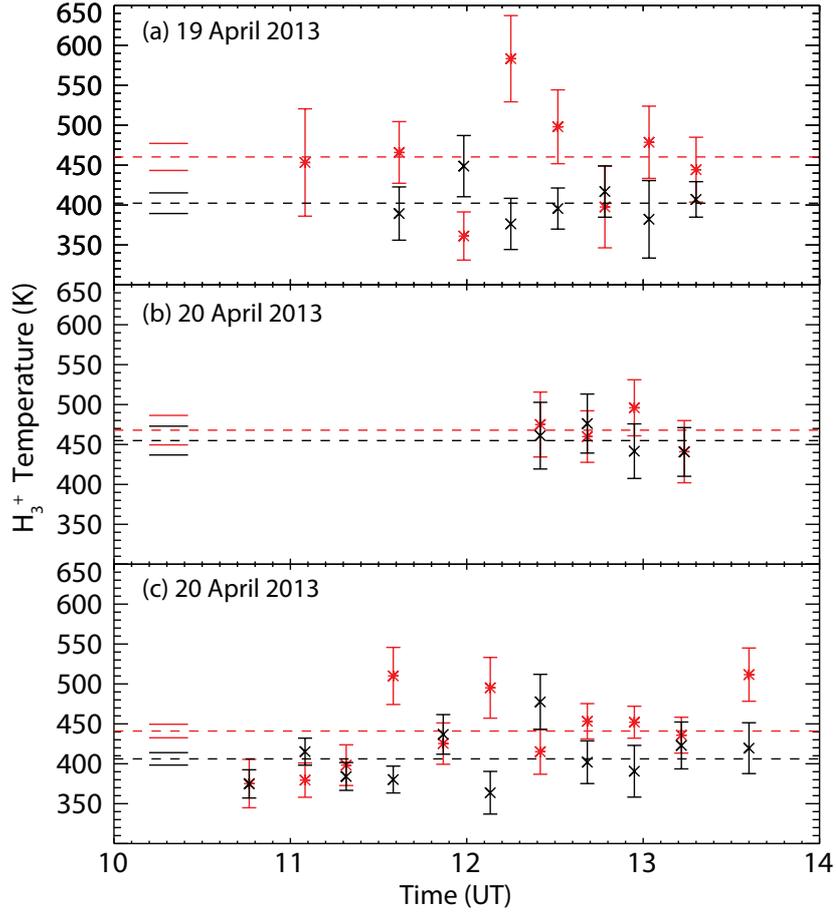


Fig. 6. NMO and SMO H_3^+ temperature as a function of observation time. The three panels show H_3^+ temperatures as a function of time for the three nights of observations as indicated. The NMO values are shown as the black crosses, while the SMO values are shown as the red asterisks. The uncertainties listed are 1-sigma and arise from the S/N of the spectral fit. Note that the northern main oval temperature of 994 ± 900 K (in the first row of Table 2) is not shown in panel (a), as it is assumed to be unphysical (this was possibly due to a passing cirrus cloud, reducing the S/N). The black and red dashed horizontal lines show the average temperature of all the plotted data points for north and south, respectively, with associated 1-sigma uncertainties above and below shown as short solid lines.

384 magnetically conjugate location in the other [*Meredith et al.* 2014].

385 The total emission shown in Tables 2 - 4 and Figure 8 is higher in the NMO for nearly all
 386 data points compared to the SMO - a similar trend is seen in column density, but in this case

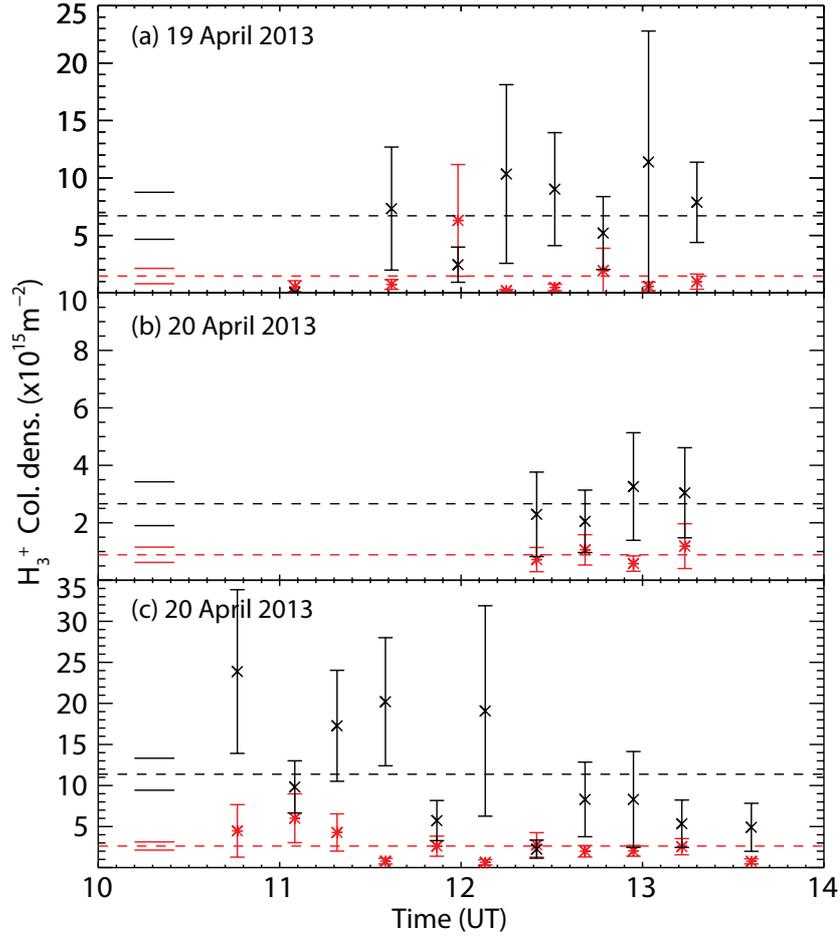


Fig. 7. NMO and SMO H_3^+ column densities as a function of time. The figure format is the same as Figure 6.

387 with smaller uncertainties. The total emission is a direct measure of H_3^+ cooling to space, so it
 388 might be argued that in the NMO, the larger quantity of H_3^+ would have led to a higher rate of
 389 thermospheric cooling, which in turn has led to lower temperatures. However, the observations
 390 by *O'Donoghue et al.* [2014] are a counter example in that high densities are associated with low
 391 total emissions, so this is not an obvious cause. Furthermore, the global circulation modeling
 392 (GCM) results of *Mueller-Wodarg et al.* [2012] of Saturn during equinoctial conditions show
 393 that H_3^+ acts only as a minor coolant in the thermosphere. The major heating mechanism
 394 in the auroral thermosphere is Joule heating, whilst adiabatic cooling and advection are the
 395 major heat sinks in the upper polar atmosphere. The densities observed here are similar to
 396 *O'Donoghue et al.* [2014] and are within the 1 to $12 \times 10^{15} \text{m}^{-2}$ range of values modeled by

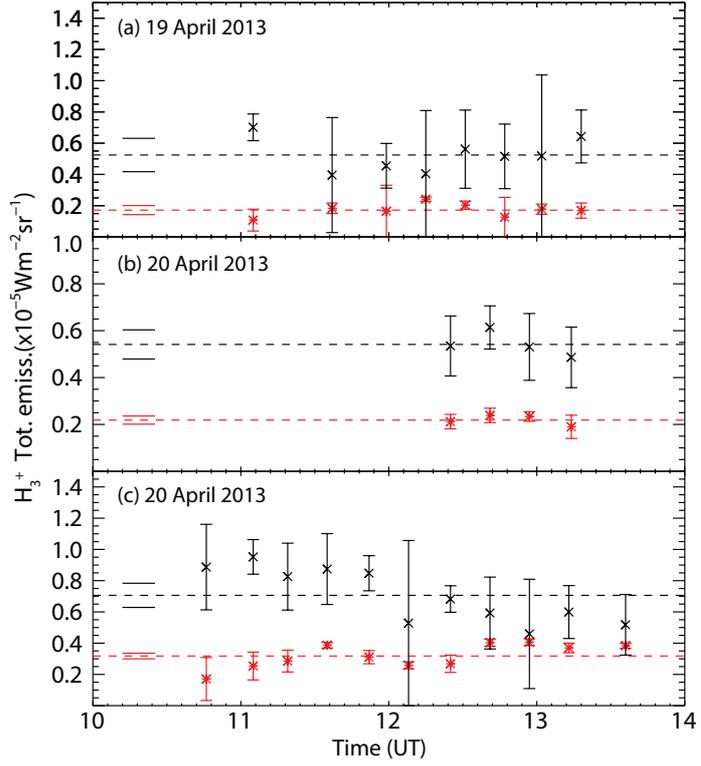


Fig. 8. NMO and SMO H_3^+ total emission as a function of time. The figure format is the same as Figure 6.

397 *Mueller-Wodarg et al.* [2012]. There are no obvious trends found here that lead us to conclude
 398 a dependence of H_3^+ parameters with system III CML. The NMO and SMO individually show
 399 sporadic variability of several 10's of K throughout all CML's, indicating little or no observable
 400 relationship.

402 In Figure 9 we show the H_3^+ parameters of all of the four previously mentioned spatial
403 regions (as shown in Figure 3) as a function of system III longitude (CML). Before continuing
404 we note that the nearby components of the north are close together and therefore subject to
405 latitudinal smearing, i.e. cross-contamination, even though gaps were left between the target
406 areas. This is due to (mainly) atmospheric scintillation/seeing effects and telescope movement
407 during spectral image exposures. However, comparison between the northern main oval and
408 midnight are separated significantly enough that these effects are negligible. First, we find that
409 there are no obvious trends leading us to conclude a dependence of H_3^+ parameters with CML.
410 The northern and southern main ovals individually show sporadic variability of several 10's of
411 K throughout all CMLs, indicating little or no observable relationship. However, the northern
412 main oval (black crosses) total emission and column density do appear to have significantly
413 higher values than the average near 50-100° CML, and this will be discussed in the next section.
414 A lack of an obvious pattern is perhaps unsurprising as there are no known CML dependencies
415 of Saturn's magnetic field. Our interests here therefore lie mainly in the average behavior of
416 each region from the combined three days of observations. The CMLs for the northern midnight
417 main oval are shifted by 180 degrees as they are on the 'night' (but sunlit) side of the planet,
418 whilst the northern polar cap (which straddles both sides) uses northern main oval CMLs. The
419 effective column integrated H_3^+ temperature is on average 465 K at midnight, 53 K greater than
420 in the main oval. Column density averages are $1 \times 10^{15} \text{ m}^{-2}$ at midnight and $8.6 \times 10^{15} \text{ m}^{-2}$ at
421 noon, similar to values produced through modeling efforts by *Moore et al.* [2004], though these
422 were produced by solar EUV alone (i.e., non-auroral conditions). Finally, the total emission
423 is $0.6 \times 10^{-5} \text{ Wm}^{-2}\text{sr}^{-1}$ at noon and $0.18 \times 10^{-5} \text{ Wm}^{-2}\text{sr}^{-1}$ at northern midnight. The polar
424 aurora temperature is 439 K on average, whilst the column density and total emission values
425 are 45% and 75%, respectively, of the northern main oval values, indicating that perhaps this

426 region is contaminated by its neighbors through the seeing effects mentioned above. Southern
 427 parameters have already been discussed in the context of their northern counterparts, but
 428 appear to be most similar to the northern midnight main oval.

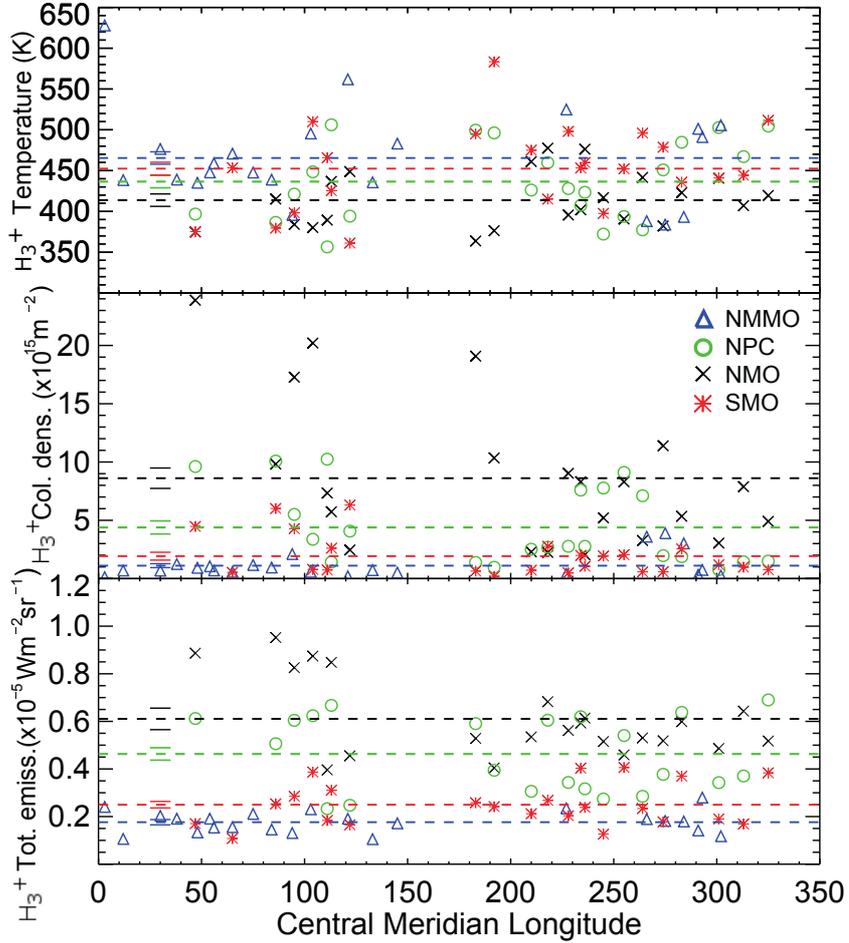


Fig. 9. Northern H_3^+ properties as a function of Saturn system III CML. Here we show the northern H_3^+ temperature, column density and total emission in panels (a), (b) and (c), respectively as a function of central meridian longitude. The different regions of interest are the northern main oval (black crosses), polar cap (green circles), midnight aurorae (blue triangles) and southern main oval (red asterisks). The average values for each are shown as dashed horizontal lines with 1-sigma uncertainty bars as short solid lines above and below to the left of the figure. The northern values at $\sim 62^\circ$ CML are not shown and not included in the calculation of average values due to high uncertainties described earlier.

429 *Mueller-Wodarg et al.* [2012] modeled Saturn's upper atmosphere for equinoctial conditions,
430 including the effects of solar radiation, magnetospheric electron precipitation and the contribu-
431 tion to the total heating rate provided by Joule heating and ion drag. The authors calculated
432 auroral H_3^+ temperatures (at 78° southern latitude) of ~ 419 K at midnight, 1 - 2 K warmer
433 than at noon. Although these temperatures are similar in absolute terms to those observed in
434 this work, the difference between the noon and midnight sectors is clearly much greater here
435 (55 K); the reason for this midnight temperature enhancement is unknown. The column den-
436 sity, on the other hand, was modeled to be $\sim 12 \times 10^{15} \text{ m}^{-2}$ at noon, compared with $\sim 1 \times 10^{15}$
437 m^{-2} at midnight, similar to that observed here. The northern column emission is a factor of
438 ~ 3 higher at noon compared to midnight in our observations, yet a factor of 15 different in the
439 above model. There are thus some areas of agreement between the model of *Mueller-Wodarg*
440 *et al.* [2012] and the observations presented here, though the relative noon-midnight differences
441 between parameters are quite large. Cross-contamination between the polar cap and the main
442 oval due to atmospheric seeing may play a role in the observation-model factor differences
443 between noon and midnight. The higher noon density and emission is likely to be driven by
444 the higher levels of 10 keV electron flux there, in accordance with the predicted maximum
445 flux at 08:00 Saturn local time (SLT), which then diminishes to a minimum near midnight
446 [*Lamy et al.* 2009]. The parameters obtained in the polar region shown by the green circles in
447 Figure 9 appear essentially to be the average of the other northern values. The activity here
448 could be maintained by transport from the midnight and noon sectors, as well as be modulated
449 by particle precipitation along open field lines which connect the planet directly to the solar
450 wind.

452 In the last section, although there was no clear organisation with CML, there were a number
453 of high density and emission values in the northern main oval at around 50-100° CML in Figure
454 9. In addition, this is a region in which we have a complete view of the 8 - 20° co-latitudes that
455 define it (compared with the limited southern main oval field of view of 18 - 22°), so it is an
456 ideal place to explore any short-term variability; in particular, that imposed by the planetary
457 period oscillations of the magnetic field. In the four panels of Figure 10 we plot each of the
458 NMO H_3^+ parameters from all three days as a function of PPO phase, Ψ_N , between 0° and
459 360°. In Figure 10 panel (a) we plot the H_3^+ Q(1,0) line intensity versus the northern PPO
460 phase, and we find a factor of ~ 2 higher intensity between 90 - 135°. The line intensity is a
461 useful metric for the overall activity of H_3^+ as it is directly observed and is a function of both
462 temperature and density. The location of the center of the fitted Gaussian distribution curve
463 (the peak) shown over-plotted in black is located at 115° and has a FWHM of 44°. Figure 10
464 panel (b) shows the NMO temperature against northern PPO phase, and this anti-correlates
465 with the column density shown in panel (c) with a Spearman's rank coefficient $r = -0.95$
466 (with a probability that these values are uncorrelated of $p < 0.01$). This and other correlations
467 between H_3^+ parameters are given in Table 6. The column density Gaussian curve peaks at
468 118° and has a FWHM of 49°- almost identical in location to the Q(1,0) line peak. In panel
469 (d) the total emission the Gaussian curve peaks at 114° with a FWHM of 40°.

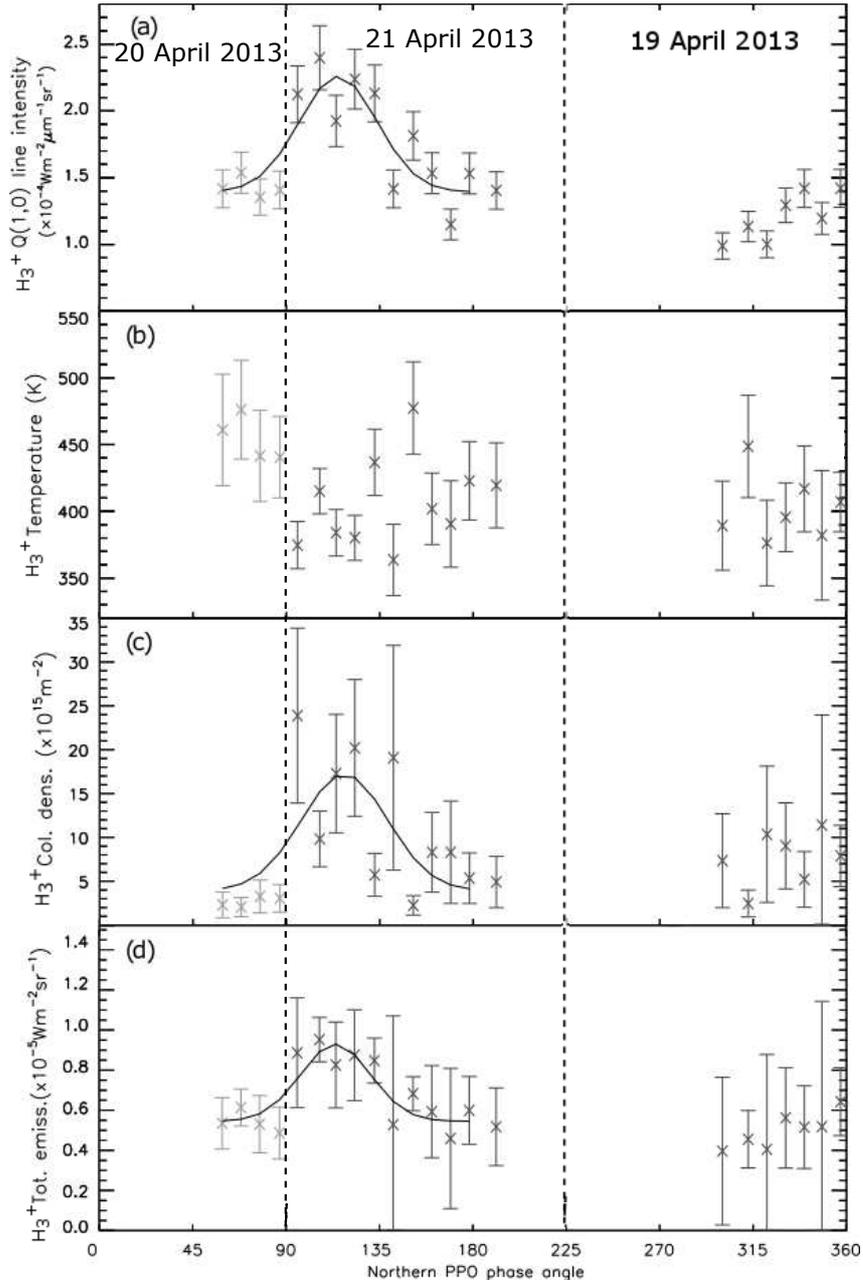


Fig. 10. NMO H_3^+ parameters as a function of northern PPO phase. Here we show the northern main oval results from the three days of this study as a function of the PPO phase angle described in the main text. The following H_3^+ parameters are shown in each of the four panels: (a) Q(1,0) line intensity, (b) temperature, (c) column density and (d) total emission.

470 The theoretical peak particle precipitation is thought to occur at $\Psi_N = 90^\circ$ as discussed
 471 in the introduction, so the above locations are some 25 degrees later on in PPO phase (1
 472 hour and 40 minutes earlier in Saturn LT). First we note that the phase model is accurate to

H ₃ ⁺ parameter	Temperature	Column density	Total emission
Q(1,0) intensity	$r = -0.04$ ($p = 0.85$)	$r = 0.25$ ($p = 0.23$)	$r = -0.79$ ($p < 0.01$)
Total emission	$r = 0.08$ ($p = 0.73$)	$r = 0.17$ ($p = 0.43$)	-
Column density	$r = -0.95$ ($p < 0.01$)	-	-

Table 6

Spearman’s rank correlation coefficients between H₃⁺ parameters when arranged in order of PPO phase.

473 approximately $\pm 10^\circ$, so this may account for some of the deviation from expectations [*Provan*
474 *et al.* 2014]. Second, the FWHM is approximately 44° for the peaks above, a considerable spread
475 in longitude; a reason for this may be the fact that our measurements are based on spectral
476 image exposures that are ~ 15 minutes in length and thus accuracy is limited to approximately
477 $\pm 5^\circ$ in PPO phase. Finally, the position of the starting location/time of the measured peak
478 in density and emission could be shifted forward due to the chemical lifetime of H₃⁺ being
479 approximately 100 - 1000 seconds [*Badman et al.* 2014]. The lifetime of H₃⁺ is also likely to
480 extend the end location/time of the Gaussian profile. Here by combining the recombination
481 rate $17.32 \times 10^{-7} \text{cm}^3 \text{s}^{-1}$ from *Moses and Bass* [2000] with typical values for the temperature
482 and number density in the auroral region at these altitudes, 450 K and $1 \times 10^4 \text{cm}^{-3}$ [*Mueller-*
483 *Wodarg et al.* 2012], we obtain an H₃⁺ lifetime of ~ 1230 seconds. These factors when combined
484 could result in the Gaussian profile being shifted by up to 20 degrees CML/phase angle later,
485 so the results herein are not inconsistent with the predicted periodic enhancement in electron
486 influx.

487 We have indicated the results from different days in the panels of Figure 10. The majority
488 of the curvature of the profile coming from the data taken on 21 April. As this dataset has
489 no overlapping PPO phase data from the different days, we cannot rule out that the observed
490 patterns are due to an enhancement in particle precipitation driven by other mechanisms.
491 For example, an interplanetary magnetic field (IMF) pointing northward can lead to magnetic

492 reconnection at low latitudes, such that planetary field lines become open and connect with
493 the solar wind [*Badman et al.* 2013]. A combination of longer observations and overlapping
494 data over the same PPO phases are required in order to definitively confirm the findings here.

495 Interestingly, the temperature appears to be lowest where the influx of charged particles is
496 highest. This could be in part due to a slight cooling effect of H_3^+ whereby it radiates heat to
497 space, although modeling work has shown such cooling is minor compared to other processes
498 like adiabatic cooling [*Mueller-Wodarg et al.* 2012]. Given the uncertainties in column density,
499 it is possible that the anti-correlations are not entirely physical and are tainted by the least-
500 squares fitting routine employed herein [*Melin et al.* 2014]. However, the trends in Figure
501 10 are arrived at independently from the fitting routine in panel (a) and through a combined
502 temperature and column density in panel (d), thus we have shown multiple instances of possible
503 H_3^+ -PPO phase dependence.

504 Analysis of the other regions (SMO, NPC, NMMO) did not yield similar correlations (or at
505 least, not as strongly) to that of the NMO, although those are regions of lower spatial resolution
506 and higher cross-latitude contamination due to seeing effects. Given the significant variability
507 seen here, it is important that similar future research include the contributions made by the
508 PPO perturbation.

509 **5 Summary**

510 On April 19, 20 and 21, the ground-based Keck telescope was employed to simultaneously
511 measure H_3^+ parameters (temperature, density and total emission) in four specific regions of
512 Saturn's ionosphere/thermosphere: 1) the northern noon region of the main auroral oval; 2) the
513 northern midnight main oval; 3) the northern polar cap and 4) the southern noon main oval. In
514 these locations, the 115 minutes of captured exposures on Saturn were used to derive ninety-two

515 H_3^+ temperatures, column densities and total emissions spread over timescales of both hours and
516 days, and therefore over a wide range of Saturn system III longitudes (CMLs) and planetary
517 period oscillation (PPO) phase angles. We have found that column integrated thermospheric
518 temperatures in the northern main oval are cooler than their southern counterparts by tens
519 of K on average. Although the northern aurorae is at times hotter than the south for some
520 individual measurements, this work lends support the hypothesis that the total thermospheric
521 heating rate (Joule heating and ion drag) is inversely proportional to magnetic field strength,
522 as discussed by *O'Donoghue et al.* [2014]. The midnight portion of the oval is on average 55 K
523 warmer than it is at noon, but the cause for this is unclear. The main oval column integrated H_3^+
524 density and emission is lower at northern midnight than it is at noon, in agreement with a peak
525 in the electron influx at 08:00 Saturn local time and a minimum flux at midnight. When the
526 northern main oval parameters of H_3^+ are ordered into the northern PPO phase we see a large
527 peak in H_3^+ density and emission at $\sim 115^\circ$ northern phase, with a full-width at half-maximum
528 (FWHM) of $\sim 44^\circ$. We find that these peaks are most likely due to the expected theoretical
529 enhancement in the influx of electrons associated with the PPO phase at 90° . A combination
530 of the H_3^+ reaction time to the influx due to ion chemistry timescales, the $\pm 10^\circ$ uncertainty in
531 the location of a given PPO phase and the lifetime of H_3^+ are likely to be partly responsible
532 for the observed peaks in H_3^+ density and emission occurring later in time (forward in phase)
533 of the expected precipitation location.

534 **Acknowledgements** The data presented herein were obtained at the W.M. Keck Observa-
535 tory, which is operated as a scientific partnership among the California Institute of Technology,
536 the University of California, and NASA. We are particularly grateful to the observing staff in
537 both Waimea and Mauna Kea for their kind assistance and we praise their ability to seem-
538 ingly clear the sky of clouds whenever we observe. The observations were made to support the
539 Cassini auroral campaign in April 2013. Discussions within the international team lead by Tom
540 Stallard on ‘Comparative Jovian Aeronomy’ have greatly benefited this work; this was hosted

541 by the International Space Science Institute (ISSI). The UK Science and Technology Facili-
542 ties Council (STFC) supported this work through the Studentship Enhancement Programme
543 (STEP) for J.O'D. and consolidated grant support for T.S.S., S.W.H.C. and H.M., whilst
544 S.V.B. was supported by a Royal Astronomical Society Research Fellowship. This material
545 is based upon work supported by the National Aeronautics and Space Administration under
546 Grant No. 9500303356 issued through the Planetary Astronomy Program for L.M. and J.O'D.
547 We thank the NASA Planetary Data System (PDS) for planetary parameter and viewing
548 geometry data.

549 References

- 550 Andrews, D. J., E. J. Bunce, S. W. H. Cowley, M. K. Dougherty, G. Provan, and D. J. South-
551 wood (2008), Planetary period oscillations in Saturn's magnetosphere: Phase relation of
552 equatorial magnetic field oscillations and Saturn kilometric radiation modulation, *J. Geo-*
553 *phys. Res. (Space Physics)*, *113*, A09205, doi:10.1029/2007JA012937.
- 554 Andrews, D. J., A. J. Coates, S. W. H. Cowley, M. K. Dougherty, L. Lamy, G. Provan, and
555 P. Zarka (2010), Magnetospheric period oscillations at Saturn: Comparison of equatorial
556 and high-latitude magnetic field periods with north and south Saturn kilometric radiation
557 periods, *J. Geophys. Res. (Space Physics)*, *115*, A12252, doi:10.1029/2010JA015666.
- 558 Badman, S. V., S. W. H. Cowley, J.-C. Gérard, and D. Grodent (2006), A statistical analysis
559 of the location and width of Saturn's southern auroras, *Ann. Geophys.*, *24*, 3533–3545, doi:
560 10.5194/angeo-24-3533-2006.
- 561 Badman, S. V., D. J. Andrews, S. W. H. Cowley, L. Lamy, G. Provan, C. Tao, S. Kasahara,
562 T. Kimura, M. Fujimoto, H. Melin, T. Stallard, R. H. Brown, and K. H. Baines (2012a),
563 Rotational modulation and local time dependence of Saturn's infrared H_3^+ auroral intensity,
564 *J. Geophys. Res.*, *117*, A09228, doi:10.1029/2012JA017990.
- 565 Badman, S. V., N. Achilleos, C. S. Arridge, K. H. Baines, R. H. Brown, E. J. Bunce, A. J.

566 Coates, S. W. H. Cowley, M. K. Dougherty, M. Fujimoto, G. Hospodarsky, S. Kasahara,
567 T. Kimura, H. Melin, D. G. Mitchell, T. Stallard, and C. Tao (2012b), Cassini observations of
568 ion and electron beams at Saturn and their relationship to infrared auroral arcs, *J. Geophys.*
569 *Res.*, *117*, A01211, doi:10.1029/2011JA017222.

570 Badman, S. V., A. Masters, H. Hasegawa, M. Fujimoto, A. Radioti, D. Grodent, N. Sergis,
571 M. K. Dougherty, and A. J. Coates (2013), Bursty magnetic reconnection at Saturn's mag-
572 netopause, *Geophys. Res. Lett.*, *40*, 1027–1031, doi:10.1002/grl.50199.

573 Badman, S. V., G. Branduardi-Raymont, M. Galand, S. L. G. Hess, N. Krupp, L. Lamy,
574 H. Melin, and C. Tao (2014), Auroral Processes at the Giant Planets: Energy Deposition,
575 Emission Mechanisms, Morphology and Spectra, *Space Sci. Rev.*, doi:10.1007/s11214-014-
576 0042-x.

577 Barrow, D. J., and K. I. Matcheva (2013), Modeling the effect of atmospheric gravity waves
578 on Saturn's ionosphere, *Icarus*, *224*, 32–42, doi:10.1016/j.icarus.2013.01.027.

579 Belenkaya, E. S., S. W. H. Cowley, C. J. Meredith, J. D. Nichols, V. V. Kalegaev, I. I. Alexeev,
580 O. G. Barinov, W. O. Barinova, and M. S. Blokhina (2014), Magnetospheric magnetic field
581 modelling for the 2011 and 2012 HST Saturn aurora campaigns - implications for auroral
582 source regions, *Ann. Geophys.*, *32*, 689–704, doi:10.5194/angeo-32-689-2014.

583 Brown, R. H., K. H. Baines, G. Bellucci, J.-P. Bibring, B. J. Buratti, F. Capaccioni, P. Cer-
584 roni, R. N. Clark, A. Coradini, D. P. Cruikshank, P. Drossart, V. Formisano, R. Jaumann,
585 Y. Langevin, D. L. Matson, T. B. McCord, V. Mennella, E. Miller, R. M. Nelson, P. D.
586 Nicholson, B. Sicardy, and C. Sotin (2004), The Cassini Visual And Infrared Mapping Spec-
587 trometer (Vims) Investigation, *Space Science Reviews*, *115*, 111–168, doi:10.1007/s11214-
588 004-1453-x.

589 Bunce, E. J., C. S. Arridge, J. T. Clarke, A. J. Coates, S. W. H. Cowley, M. K. Dougherty,
590 J.-C. Gérard, D. Grodent, K. C. Hansen, J. D. Nichols, D. J. Southwood, and D. L. Talboys
591 (2008), Origin of Saturn's aurora: Simultaneous observations by Cassini and the Hubble
592 Space Telescope, *J. Geophys. Res. (Space Physics)*, *113*, A09209, doi:10.1029/2008JA013257.

593 Burton, M. E., M. K. Dougherty, and C. T. Russell (2010), Saturn’s internal planetary magnetic
594 field, *J. Geophys. Res.*, *37*, L24105, doi:10.1029/2010GL045148.

595 Carbary, J. F. (2012), The morphology of Saturn’s ultraviolet aurora, *J. Geophys. Res. (Space*
596 *Physics)*, *117*, A06210, doi:10.1029/2012JA017670.

597 Cowley, S. W. H., and E. J. Bunce (2003), Corotation-driven magnetosphere-ionosphere cou-
598 pling currents in Saturn’s magnetosphere and their relation to the auroras, *Annales Geo-*
599 *physicae*, *21*, 1691–1707, doi:10.5194/angeo-21-1691-2003.

600 Cowley, S. W. H., E. J. Bunce, and J. M. O’Rourke (2004), A simple quantitative model
601 of plasma flows and currents in Saturn’s polar ionosphere, *J. Geophys. Res.*, *109*, A05212,
602 doi:10.1029/2003JA010375.

603 Grodent, D., J. H. Waite, Jr., and J.-C. Gérard (2001), A self-consistent model of the Jovian
604 auroral thermal structure, *J. Geophys. Res.*, *106*, 12,933–12,952, doi:10.1029/2000JA900129.

605 Grodent, D., J. Gustin, J.-C. Gérard, A. Radioti, B. Bonfond, and W. R. Pryor (2011),
606 Small-scale structures in Saturn’s ultraviolet aurora, *J. Geophys. Res. (Space Physics)*, *116*,
607 A09225, doi:10.1029/2011JA016818.

608 Gurnett, D. A., A. Lecacheux, W. S. Kurth, A. M. Persoon, J. B. Groene, L. Lamy, P. Zarka,
609 and J. F. Carbary (2009), Discovery of a north-south asymmetry in Saturn’s radio rotation
610 period, *Geophys. Res. Lett.*, *36*, L16102, doi:10.1029/2009GL039621.

611 Hunt, G. J., S. W. H. Cowley, G. Provan, E. J. Bunce, I. I. Alexeev, E. S. Belenkaya, V. V. Kale-
612 gaev, M. K. Dougherty, and A. J. Coates (2014), Field-aligned currents in Saturn’s southern
613 nightside magnetosphere: Sub-corotation and planetary period oscillation components., *J.*
614 *Geophys. Res. (submitted)*.

615 Kaiser, M. L., M. D. Desch, J. W. Warwick, and J. B. Pearce (1980), Voyager de-
616 tecton of nonthermal radio emission from Saturn, *Science*, *209*, 1238–1240, doi:
617 10.1126/science.209.4462.1238.

618 Kao, L., T. Oka, S. Miller, and J. Tennyson (1991), A table of astronomically important ro-
619 vibrational transitions for the H₃(+) molecular ion, *ApJS*, *77*, 317–329, doi:10.1086/191606.

620 Koskinen, T. T., B. R. Sandel, R. V. Yelle, F. J. Capalbo, G. M. Holsclaw, W. E. Mc-
621 Clintock, and S. Edgington (2013), The density and temperature structure near the
622 exobase of Saturn from Cassini UVIS solar occultations, *Icarus*, *226*, 1318–1330, doi:
623 10.1016/j.icarus.2013.07.037.

624 Lam, H. A., N. Achilleos, S. Miller, J. Tennyson, L. M. Trafton, T. R. Geballe, and G. E.
625 Ballester (1997), A Baseline Spectroscopic Study of the Infrared Auroras of Jupiter, *Icarus*,
626 *127*, 379–393, doi:10.1006/icar.1997.5698.

627 Lamy, L., B. Cecconi, R. Prangé, P. Zarka, J. D. Nichols, and J. T. Clarke (2009), An auroral
628 oval at the footprint of Saturn’s kilometric radio sources, colocated with the UV aurorae, *J.*
629 *Geophys. Res. (Space Physics)*, *114*, A10212, doi:10.1029/2009JA014401.

630 Lamy, L., R. Prangé, W. Pryor, J. Gustin, S. V. Badman, H. Melin, T. Stallard, D.-G.
631 Mitchell, and P. C. Brandt (2013), Multispectral simultaneous diagnosis of Saturn’s au-
632 rorae throughout a planetary rotation, *J. Geophys. Res. (Space Physics)*, *118*, 4817–4843,
633 doi:10.1002/jgra.50404.

634 McLean, I. S., E. E. Becklin, O. Bendiksen, G. Brims, J. Canfield, D. F. Figer, J. R. Graham,
635 J. Hare, F. Lacayanga, J. E. Larkin, S. B. Larson, N. Levenson, N. Magnone, H. Teplitz, and
636 W. Wong (1998), Design and development of NIRSPEC: a near-infrared echelle spectrograph
637 for the Keck II telescope, in *Society of Photo-Optical Instrumentation Engineers (SPIE)*
638 *Conference Series, Society of Photo-Optical Instrumentation Engineers (SPIE) Conference*
639 *Series*, vol. 3354, edited by A. M. Fowler, pp. 566–578.

640 Melin, H. (2006), Comparative Aeronomy of the Upper Atmospheres of the Giant Planets,
641 *University College London*.

642 Melin, H., S. Miller, T. Stallard, L. M. Trafton, and T. R. Geballe (2007), Variability in the H_3^+
643 emission of Saturn: Consequences for ionisation rates and temperature, *Icarus*, *186*, 234–241,
644 doi:10.1016/j.icarus.2006.08.014.

645 Melin, H., T. Stallard, S. Miller, J. Gustin, M. Galand, S. V. Badman, W. R. Pryor,
646 J. O’Donoghue, R. H. Brown, and K. H. Baines (2011), Simultaneous Cassini VIMS

647 and UVIS observations of Saturn’s southern aurora: Comparing emissions from H, H₂
648 and H₃⁺ at a high spatial resolution, *Geophysical Research Letters*, *38*, L15203, doi:
649 10.1029/2011GL048457.

650 Melin, H., T. S. Stallard, J. O’Donoghue, S. V. Badman, S. Miller, and J. S. D. Blake (2014),
651 On the anticorrelation between H₃⁺ temperature and density in giant planet ionospheres,
652 *Mon. Not. R. Astron. Soc.*, *438*, 1611–1617, doi:10.1093/mnras/stt2299.

653 Meredith, C., S. W. H. Cowley, K. C. Hansen, J. D. Nichols, and T. K. Yeoman (2013),
654 Simultaneous conjugate observations of small-scale structures in Saturn’s dayside ultraviolet
655 auroras - implications for physical origins, *J. Geophys. Res. (Submitted)*.

656 Meredith, C. J., I. I. Alexeev, S. V. Badman, E. S. Belenkaya, S. W. H. Cowley, M. K.
657 Dougherty, V. V. Kalegaev, G. R. Lewis, and J. D. Nichols (2014), Saturn’s dayside
658 ultraviolet auroras: Evidence for morphological dependence on the direction of the up-
659 stream interplanetary magnetic field, *J. Geophys. Res. (Space Physics)*, *119*, 1994–2008,
660 doi:10.1002/2013JA019598.

661 Miller, S., R. D. Joseph, and J. Tennyson (1990), Infrared emissions of H₃(+) in the atmosphere
662 of Jupiter in the 2.1 and 4.0 micron region, *Astrophys. J.*, *360*, L55–L58, doi:10.1086/185811.

663 Miller, S., T. Stallard, C. Smith, and et al. (2006), H₃⁺: the driver of giant planet atmospheres,
664 *Phil. Trans. Roy. Soc. London.*, *364*, 3121–3137, doi:10.1098/rsta.2006.1877.

665 Miller, S., T. Stallard, H. Melin, and J. Tennyson (2010), H₃⁺ cooling in planetary atmospheres,
666 *Faraday Discussions*, *147*, 283, doi:10.1039/c004152c.

667 Moore, L., M. Galand, I. Mueller-Wodarg, R. Yelle, and M. Mendillo (2008), Plasma tem-
668 peratures in Saturn’s ionosphere, *J. Geophys. Res. (Space Physics)*, *113*, A10306, doi:
669 10.1029/2008JA013373.

670 Moore, L. E., M. Mendillo, I. C. F. Müller-Wodarg, and D. L. Murr (2004), Modeling of
671 global variations and ring shadowing in Saturn’s ionosphere, *Icarus*, *172*, 503–520, doi:
672 10.1016/j.icarus.2004.07.007.

673 Moses, J. I., and S. F. Bass (2000), The effects of external material on the chemistry and struc-

674 ture of Saturn’s ionosphere, *J. Geophys. Res.*, *105*, 7013–7052, doi:10.1029/1999JE001172.

675 Mueller-Wodarg, I. C. F., L. Moore, M. Galand, S. Miller, and M. Mendillo (2012),

676 Magnetosphere-atmosphere coupling at Saturn: 1 - Response of thermosphere and iono-

677 sphere to steady state polar forcing, *Icarus*, *221*, 481–494, doi:10.1016/j.icarus.2012.08.034.

678 Neale, L., S. Miller, and J. Tennyson (1996), Spectroscopic Properties of the H_3^+ Molecule: A

679 New Calculated Line List, *The Astrophysical Journal*, *464*, 516, doi:10.1086/177341.

680 Nichols, J. D., S. V. Badman, E. J. Bunce, J. T. Clarke, S. W. H. Cowley, F. J. Crary, M. K.

681 Dougherty, J.-C. Gérard, D. Grodent, K. C. Hansen, W. S. Kurth, D. G. Mitchell, W. R.

682 Pryor, T. S. Stallard, D. L. Talboys, and S. Wannawichian (2009), Saturn’s equinoctial

683 auroras, *Geophysical Research Letters*, *36*, L24102, doi:10.1029/2009GL041491.

684 O’Donoghue, J., T. S. Stallard, H. Melin, G. H. Jones, S. W. H. Cowley, S. Miller, K. H.

685 Baines, and J. S. D. Blake (2013), The domination of Saturn’s low-latitude ionosphere by

686 ring ‘rain’, *Nature*, *496*, 193–195, doi:10.1038/nature12049.

687 O’Donoghue, J., T. S. Stallard, H. Melin, S. W. H. Cowley, S. V. Badman, L. Moore, S. Miller,

688 C. Tao, K. H. Baines, and J. S. D. Blake (2014), Conjugate observations of Saturn’s northern

689 and southern H_3^+ aurorae, *Icarus*, *229*, 214–220, doi:10.1016/j.icarus.2013.11.009.

690 Oka, T. (2006), Interstellar Chemistry Special Feature: Interstellar H_3^+ , *Proceedings of the*

691 *National Academy of Science*, *103*, 12,235–12,242, doi:10.1073/pnas.0601242103.

692 Provan, G., D. J. Andrews, C. S. Arridge, A. J. Coates, S. W. H. Cowley, S. E. Milan, M. K.

693 Dougherty, and D. M. Wright (2009), Polarization and phase of planetary-period magnetic

694 field oscillations on high-latitude field lines in Saturn’s magnetosphere, *J. Geophys. Res.*

695 *(Space Physics)*, *114*, A02225, doi:10.1029/2008JA013782.

696 Provan, G., D. J. Andrews, C. S. Arridge, A. J. Coates, S. W. H. Cowley, G. Cox, M. K.

697 Dougherty, and C. M. Jackman (2012), Dual periodicities in planetary-period magnetic

698 field oscillations in Saturn’s tail, *J. Geophys. Res. (Space Physics)*, *117*, A01209, doi:

699 10.1029/2011JA017104.

700 Provan, G., L. Lamy, S. W. H. Cowley, and M. K. Dougherty (2014), Planetary period os-

701 cillations in Saturn's magnetosphere: Comparison of magnetic oscillations and SKR mod-
702 ulations in the post-equinox interval, *J. Geophys. Res. (Space Physics)*, submitted, doi:
703 10.1002/2014JA020011.

704 Radioti, A., D. Grodent, J.-C. Gérard, S. E. Milan, B. Bonfond, J. Gustin, and W. Pryor
705 (2011), Bifurcations of the main auroral ring at Saturn: ionospheric signatures of con-
706 secutive reconnection events at the magnetopause, *J. Geophys. Res.*, *116*, A11209, doi:
707 10.1029/2011JA016661.

708 Radioti, A., D. Grodent, J.-C. Gérard, B. Bonfond, J. Gustin, W. Pryor, J. M. Jasinski, and
709 C. S. Arridge (2013), Auroral signatures of multiple magnetopause reconnection at Saturn,
710 *Geophys. Res. Lett.*, *40*, 4498–4502, doi:10.1002/grl.50889.

711 Raynaud, E., E. Lellouch, J.-P. Maillard, G. R. Gladstone, J. H. Waite, B. Bézard, P. Drossart,
712 and T. Fouchet (2004), Spectro-imaging observations of Jupiter's 2- μ m auroral emission. I.
713 H₃⁺ distribution and temperature, *Icarus*, *171*, 133–152, doi:10.1016/j.icarus.2004.04.020.

714 Sandel, B. R., and A. L. Broadfoot (1981), Morphology of Saturn's aurora, *Nature*, *292*, 679–
715 682, doi:10.1038/292679a0.

716 Sandel, B. R., D. E. Shemansky, A. L. Broadfoot, J. B. Holberg, G. R. Smith, J. C. McConnell,
717 D. F. Strobel, S. K. Atreya, T. M. Donahue, H. W. Moos, D. M. Hunten, R. B. Pomphrey,
718 and S. Linick (1982), Extreme ultraviolet observations from the Voyager 2 encounter with
719 Saturn, *Science*, *215*, 548–553, doi:10.1126/science.215.4532.548.

720 Smith, C. G. A., A. D. Aylward, G. H. Millward, S. Miller, and L. E. Moore (2007),
721 An unexpected cooling effect in Saturn's upper atmosphere, *Nature*, *445*, 399–401, doi:
722 10.1038/nature05518.

723 Southwood, D. (2011), Direct evidence of differences in magnetic rotation rate between Saturn's
724 northern and southern polar regions, *J. Geophys. Res. (Space Physics)*, *116*, A01201, doi:
725 10.1029/2010JA016070.

726 Southwood, D. J., and S. W. H. Cowley (2014), The origin of Saturn's magnetic periodicities:
727 Northern and southern current systems, *J. Geophys. Res. (Space Physics)*, *119*, 1563–1571,

728 doi:10.1002/2013JA019632.

729 Stallard, T. S., H. Melin, S. Miller, J. O' Donoghue, S. W. H. Cowley, S. Badman, A. Adriani,
730 R. H. Brown, and K. H. Baines (2012a), Temperature changes and energy inputs in giant
731 planet atmospheres: what we are learning from H_3^+ , *Phil. Trans. Roy. Soc.*, *370*, 5213–5224,
732 doi:10.1098/rsta.2012.0028.

733 Stallard, T. S., H. Melin, S. Miller, S. V. Badman, R. H. Brown, and K. H. Baines (2012b),
734 Peak emission altitude of Saturn's H_3^+ aurora, *Geophys. Res. Lett.*, *39*, L15103, doi:
735 10.1029/2012GL052806.

736 Yelle, R. V., and S. Miller (2004), *Jupiter's thermosphere and ionosphere*, pp. 185–218.